Ground Wtr Report - 1966

GROUND-WATER RESOURCES

of MARTIN COUNTY, NORTH CAROLINA

Division of Ground Water

GROUND-WATER BULLETIN NUMBER 9

NORTH CAROLINA
DEPARTMENT OF WATER RESOURCES





RALEGO NOVEMBER 1966

STATE OF NORTH CAROLINA DEPARTMENT OF WATER RESOURCES

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November 9, 1966

The Honorable Dan K. Moore Governor of North Carolina Raleigh, North Carolina

Dear Governor Moore:

I am pleased to submit Ground-Water Bulletin Number 9, "The Ground-Water Resources of Martin County, North Carolina" by Granville G. Wyrick, Geologist, U. S. Geological Survey.

This report contains the results of a detailed study of the ground-water resources made by the U. S. Geological Survey in cooperation with the Martin County Board of Commissioners and the North Carolina Department of Water Resources. It should prove to be of much value toward the economic and industrial development of the County.

Respectfully submitted,

George E. Pickett

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GROUND WATER RESOURCES OF MARTIN COUNTY NORTH CAROLINA

Ву

Granville G. Wyrick

ABSTRACT

Martin County is in the north-central part of the North Carolina Coastal Plain and is underlain by sedimentary rocks that thicken from about 400 feet in the western part of the county to about 1,000 feet in the eastern part. The sediments, ranging in age from Cretaceous to Recent, are composed chiefly of clay, sand, shell, and limestone that were deposited on crystalline basement rock. Pre-Miocene sediments strike north-northeast and dip to the southeast whereas Miocene and younger sediments are essentially horizontal.

Within the sedimentary section are seven sand, shell, or limestone beds which are the principal aquifers in the county. Water in the lower six of the aquifers occurs under artesian conditions, and in the seventh, under non-artesian conditions. The average transmissibility of the artesian aquifers ranges from 4,000 gallons per day per foot in Aquifer Number 4 to 18,000 gallons per day per foot, in Aquifer Number 3. The average daily recharge to the artesian aquifers in the county is computed as 22 million gallons per day for aquifers between 100 feet and 300 feet below land surface.

Sufficient quantities of water of chemical quality suitable for most purposes are available in the aquifers underlying Martin County. However, in many places wells yield water of unsuitable quality or insufficient quantity, largely due to the methods of well construction. Four principal problems of water supply apparently are related to this cause: (1) Water containing concentrations of chloride in excess of 250 parts per million is obtained from wells that are drilled to depths greater than 450 feet, after passing through aquifers containing water of usable quality; (2) water containing iron in concentrations greater than 0.5 parts per million is obtained from wells completed in the non-artesian aquifer, but water of better quality is available at slightly greater depths; (3) wells are finished in shell or limestone aquifers which generally produce

water with a hardness in excess of 200 parts per million; (4) the specific capacities of wells in the county on the basis of yield are below the specific capacities calculated on the basis of aquifer transmissibilities. Aquifer characteristics determined during this investigation indicate that these four problems may be minimized by drilling wells that tap aquifers containing water of acceptable chemical quality, and by the use of screen sizes and screen lengths that are adapted to specific aquifers.

INTRODUCTION

Ground water is one of the most important natural resources of the Coastal Plain of North Carolina. Ground water is also, perhaps, one of the least understood natural resources of the area. In some places, the development of adequate water supplies from wells has been a problem due to the lack of detailed geologic and hydrologic data. In other places, where wells do yield adequate quantities of water, the chemical quality of the water may be undesirable because of high concentrations of dissolved minerals.

The ground-water resources of Martin County are being developed at an increasing rate. With an increase in per capita use of water and an emphasis on industrial growth, this trend is likely to continue. All of the water used for domestic purposes in Martin County is ground water. A large portion of the water used for irrigation, stock, and industry also is ground water. With an increase in the use of ground water in Martin County there is an increasing urgency to understand the factors controlling the ground-water resources of the county so that both individual water supplies and the ground-water resources as a whole may be wisely and economically developed.

The Martin County Board of Commissioners under the direction of their Chairman, John H. Edwards, recognized the need for a study of the ground-water resources of the county, and in 1958 requested that the U. S. Geological Survey undertake a detailed ground-water investigation.

This investigation, the first detailed county-wide ground-water investigation in North Carolina, was financed by county funds that were matched by the U. S. Geological Survey over a 4-year period beginning in 1958. This report contains the results of that investigation.

The principal phases of the investigation included the following:

- 1. An inventory of most existing, particularly deep, wells to determine their location, depth, diameter, yield, and other pertinent data.
- 2. The drilling of test wells and auger holes in selected areas where sufficient information could not be obtained from existing wells.
- 3. Geologic studies to determine the character and extent of the various geologic formations as they influence and control the occurrence of ground water.

- 4. The determination of the water-transmitting and water-storage capacities of the aquifers.
- 5. The collection and chemical analysis of water samples to determine specific chemical characteristics of the ground water.
- 6. The analysis of data collected during the investigation and the preparation of this report for the Board of Commissioners.

The field work for the investigation began in December 1958. The author has worked on the project since that time. Other personnel who assisted on the project for short periods of time included John L. Snider and Stephen E. Kesler. Most of the paleontological determinations were made by Philip M. Brown.

Previous Investigations

The geology and ground-water resources of Martin County have been discussed briefly in several reports published by the U. S. Geological Survey and the North Carolina Department of Conservation and Development.

Mundorff (1945, p. 41-58) describes the general geology of the North Carolina Coastal Plain. Brown (1958, p. 34-35) gives detailed geologic and paleontologic descriptions of the material from three wells in Martin County. Brown (1959, p. 70-77) discusses the geology and occurrence of ground water in Martin County. That report, based on a reconnaissance study, includes records of the data collected from 73 wells, with a detailed log of one well, and the chemical analyses of water samples from 22 wells. Wyrick and Floyd (1961, p. 1-11) discuss the water-transmitting characteristics of the Castle Hayne Limestone in areas of discharge and non-discharge at selected sites in Martin County.

Acknowledgments

Many residents of Martin County cooperated in furnishing information about their wells during the investigation. The Hartsfield Water Company, Layne-Atlantic Company, and R. L. Magette Well Company were helpful in furnishing data from wells that they drilled during the investigation. County and municipal officials, including officials of the Williamston Water Department, the Martin County Industrial Development Commission, and Martin County Health Department were very cooperative.

The investigation was made under the direct supervision of Philip M. Brown, district geologist, and under the general supervision of

P. E. La Moreaux, former chief of the Ground Water Branch, and O. M. Hackett, chief of the Ground Water Branch, U. S. Geological Survey.

GEOGRAPHY

Martin County is in the north-central part of the Coastal Plain of North Carolina (fig. 1). It is bounded on the north by the Roanoke River, on the east by Welsh Creek, on the south by Beaufort and Pitt Counties, and on the west by Edgecombe and Halifax Counties. The county area is 481 square miles.

Williamston is the largest town and the county seat. Other municipalities include Beargrass, Everetts, Gold Point, Hamilton, Hassell, Jamesville, Oak City, Parmele, and Robersonville.

Climate

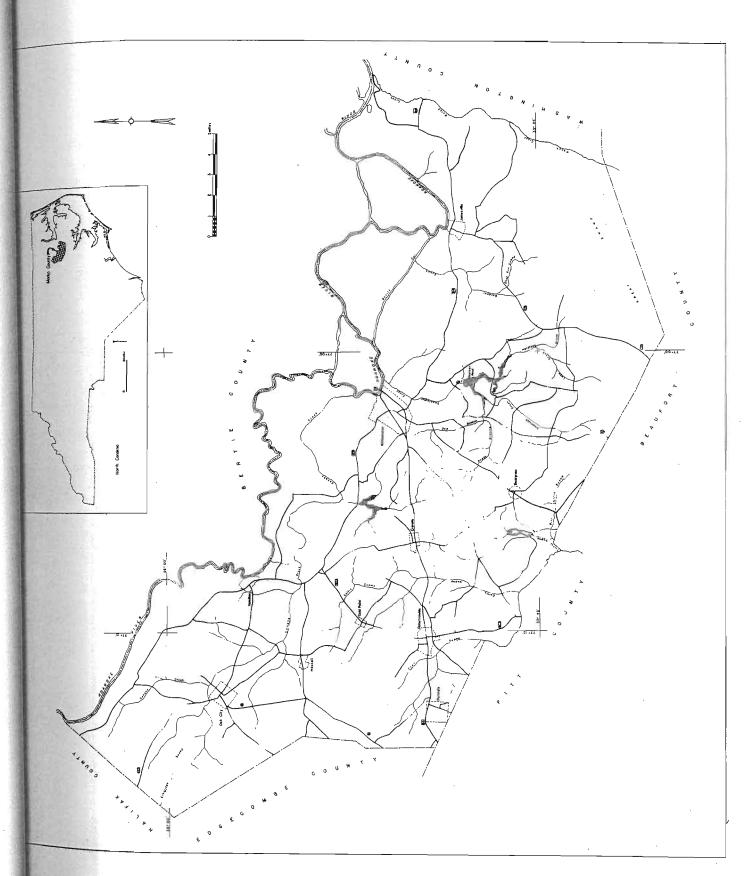
The normal average annual precipitation for Martin County during the 29-year period from 1933 to 1962 was approximately 47 inches, according to the U. S. Weather Bureau. Figure 2 shows the distribution by months, of average annual precipitation for this period. The rainfall is heaviest during the summer, but is fairly well distributed throughout the rest of the year. Records of the temperature in Martin County have been kept for nine years. The U. S. Weather Bureau does not consider this a sufficient period for calculation of the long-term mean annual temperature. However, the mean temperature for 1959 was approximately 62° F. and for 1960 it was about 61° F.

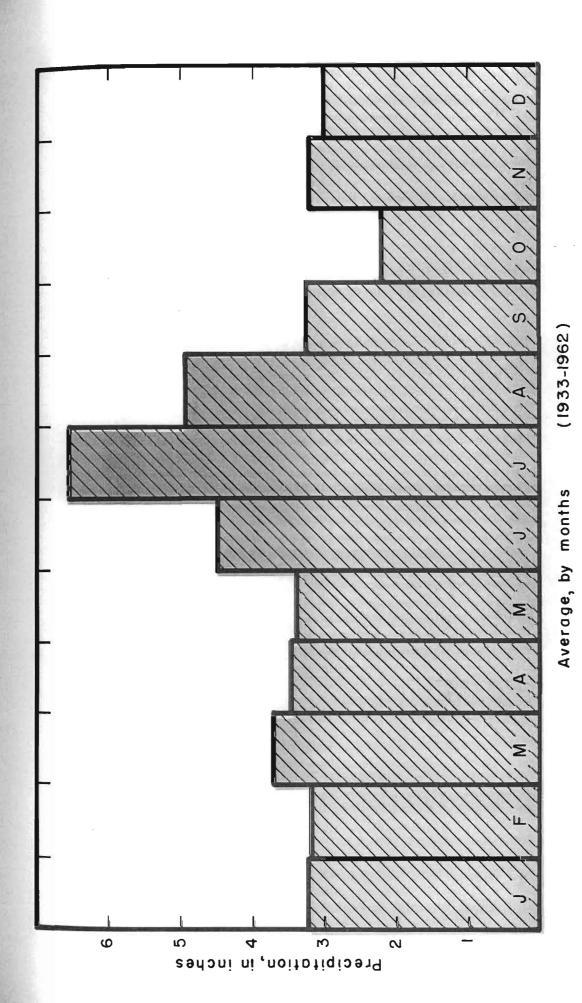
Population

Martin County had a total population of 27,139 in 1960 according to the U. S. Bureau of the Census. The census indicated that approximately 74 percent of the population is classified as farm-rural. There are approximately 4,800 homes within the county. Of these about 3,700 are rural and the owners must provide their own water supplies. On the average, one well supplies slightly more than two homes in the rural areas.

Ec on omy

The economy of Martin County is basically agricultural. The growing and processing of tobacco and peanuts account for most of the farm income. In recent years livestock raising has become increasingly important to the economy. Also, according to the Martin County Industrial





29 YEAR AVERAGE PRECIPITATION OF WILLIAMSTON FOR PERIOD OF RECORD

Development Commission, there are 24 industries in the county. For the most part, these industries are concerned with the manufacturing of wood products, textiles and textile products, and of processing food, chemicals, and pulp.

Physiography

Martin County lies within the Atlantic Coastal Plain province. Parts of the Coastal Plain have been submerged since Miocene time. Whenever the level of the sea was relatively static, in relation to adjacent land masses, for a long period, waves and currents leveled the sea floor, forming a general flat surface or plane. As land masses emerged, a plane emerged as essentially a flat land surface.

In Martin County there is evidence of three former submerged surfaces. One surface occurs at an altitude between 25 and 50 feet above mean sea level (msl), another between 50 and 75 feet above msl, and the third between 75 and 100 feet above msl (fig. 3).

The land surface in Martin County has been dissected by streams which drain north and east toward the Roanoke River and south and east toward the Pamlico River (fig. 1). In the eastern part of the county, drainage is poorly developed and large swamps occur. In the central and western parts of the county, drainage is well developed and there are few swamps.

The principal streams which drain into the Roanoke River are Conoho Creek, Sweet Water Creek, and Welch Creek. The principal streams which drain into the Pamlico River are Collie Creek and Turkey Creek. Approximately 75 percent of the county is drained by the Roanoke River.

GEOLOGY

A study of the geology of an area is essential to the study of its ground-water resources. Geologic formations of high permeability may readily store and transmit ground water, whereas, geologic formations of low permeability may retard the movements of ground water. Thus, the occurrence and movement of ground water depends, among other factors, upon the size, shape, and physical characteristics of the geologic formations. In addition, ground water may dissolve from or deposit chemicals in the material through which it moves. Thus the chemical quality of ground water is dependent upon the geology of an area.

The geology of Martin County has been studied by two principal methods. One method consisted of the collection, preparation, and examination of cuttings from wells and auger holes. The second method consisted of the collection and examination of geophysical data, including electric and gamma-ray logs from wells and test holes.

Test and Auger Holes

Geologic data were collected from 20 wells in and near Martin County. These data were supplemented by data obtained during the augering of 38 holes and the drilling of two deep test holes (fig. 4).

A correlation of the geological and geophysical data collected from eight wells is shown in figure 5. The formation boundaries were picked on the basis of microfaunal assemblages, lithology, and geophysical characteristics. The seven principal water-bearing zones are correlated on the basis of resistivity characteristics shown on electric logs and on lithology. A synthesis of all of the geological and geophysical data from the wells located on figure 4 is the basis for defining the structure of the geologic formations and the water-bearing zones in Martin County.

General Geology

Martin County is underlain by clay, silt, sand, and limestone beds which were deposited on crystalline basement rocks. These sedimentary beds generally slope to the southeast in Martin County. They were, for the most part, deposited in shallow seas. The sedimentary beds contain material that was weathered from crystalline rocks and transported by streams to the seas, mixed with marine materials such as seashells and chemical precipitates, sorted by wave and current action and deposited

LAND-SURFACE PROFILES

LOCATIONS OF GEOLOGIC AND GEOPHYSICAL DATA CONTROL WELLS

CORRELATION OF GEOLOGIC AND GEOPHYSICAL DATA

on the sea bottom, and covered by later sediments. The sequence of deposition was not continuous. Sea level fluctuated in relation to adjacent land masses. During the periods when sea level was relatively lower than adjacent land masses, the sediments, previously deposited, were above sea level, and were exposed to weathering and erosion processes. With a subsequent lowering of land masses, these eroded surfaces were inundated and more sediments were deposited upon them.

The sedimentary layers have been subdivided into formations, or other units, which consist of material of similar lithic composition, deposited during a given period of geologic time. These formations, or units, are identified by their position in the sequence of sediments, their lithic composition, and their fossil content. The surfaces that separate the formations are referred to as contacts. If there was no significant lapse between the deposition of two formations, their contact is said to be conformable; if there was a period of nondeposition or erosion between formations their contact is said to represent an unconformity. An unconformity represents a period of nondeposition or erosion and its surface is usually an undulating plane. When the layers of a formation above an unconformity are parallel to the layers of the formation below, the contact is a disconformity. When the layers of the formations above and below an unconformity are not parallel, the contact is a nonconformity. The contact relationship of the various geologic formations may greatly influence the movements of ground water between formations. These relationships (fig. 6), and the relationships of these formations to the aquifers (fig. 5), are discussed in the ground-water section.

Land surfaces adjacent to an active depositional basin slowly sink, rise, or remain static during the geologic epochs associated with the formation of the basin. When land surfaces sink, the adjacent sea encroaches on the land. This encroachment is called a transgression. When land surfaces rise, the adjacent sea recedes from the land. This recession of the sea is called a regression. In any one area, in general, fine-grained sediments are deposited during transgressions, whereas coarse-grained sediments are deposited during regressions.

Geologic Formations

The geologic formations in Martin County consist of sediments of Cretaceous age which were deposited on crystalline basement rocks, and of Tertiary and Quaternary sediments which overlie the Cretaceous deposits. Regionally, these formations form a large arc open to the east. North of Martin County the formations generally strike east; south of the county they generally strike north. Martin County's position on the arc is such that the formations within the county strike generally northeast and dip to the southeast.

The basement rocks are composed of crystalline rocks of Precambrian or early Paleozoic age; the surface cut on these rocks dips to the southeast (fig. 6).

Sediments of both Early and Late Cretaceous age unconformably overlie the basement rocks. Sediments of Early Cretaceous age are recognized in several wells in Halifax County adjacent to the northwestern corner of Martin County, and they undoubtedly extend into Martin County. Throughout the rest of the county, basement rock is overlain unconformably by the Tuscaloosa Formation of Late Cretaceous age (fig. 6).

The Tuscaloosa Formation, the lower part of the Black Creek Formation, and post-Miocene sediments were deposited during periods of regression. Parts of the Black Creek Formation, the Peedee Formation, the Beaufort Formation, and the Yorktown Formation were deposited during periods of transgression. The Castle Hayne Limestone was deposited during a static stage of basin development, when the land surface was neither sinking nor rising. Chemical precipitates predominate during static stages in this area.

Three formations of Late Cretaceous age occur in Martin County. The Tuscaloosa Formation is overlain by the Black Creek Formation which, in turn, is overlain by the Peedee Formation. The contacts between these three formations are unconformable.

The Beaufort Formation of Paleocene age overlies the Peedee Formation and, in turn, is overlain by the Castle Hayne Limestone of middle and late Eocene age. Their contact relationships also are unconformable. The base of the Yorktown Formation of late Miocene age appears as a horizontal line in figure 6 because of vertical scale exaggeration. However, since the Yorktown Formation was deposited on a highly eroded

GEOLOGICAL FORMATIONS PENETRATED BY WELLS

uneven surface, the contact between the Yorktown Formation and the older formations is not a horizontal plane (fig. 7).

The Yorktown Formation is blanketed by thin deposits of post-Miocene age. These deposits contain no diagnostic fauna but are of undoubted Quaternary age because they are part of the regional shoreline deposits laid down during static stages of a regressive cycle in Pleistocene or Recent epochs.

Each formation has distinctive structural, lithic, or paleontological characteristics that are discussed in detail in the following sections.

Basement rocks

The crystalline basement rocks are of granitic composition and contain an abundance of green mica. The upper 5 to 10 feet of the basement is weathered to a clay containing mica, feldspar, and disseminated quartz grains.

In the western part of the county, basement rock is approximately 400 feet below land surface, whereas, in the eastern part it is approximately 1,000 feet below land surface (table 1). The basement surface strikes N. 50° E. and dips about 17 feet per mile (fig. 6).

Cretaceous System

Lower Cretaceous Series

Sediments of Albian age are recognized in wells drilled in adjoining Halifax and Pitt Counties (Swain and Brown, 1963, p. 4). Because of the updip pinchout of a marine facies, caused by a structural high, Lower Cretaceous sediments, if present in Martin County, are nonmarine equivalents and have little importance as water-bearing units.

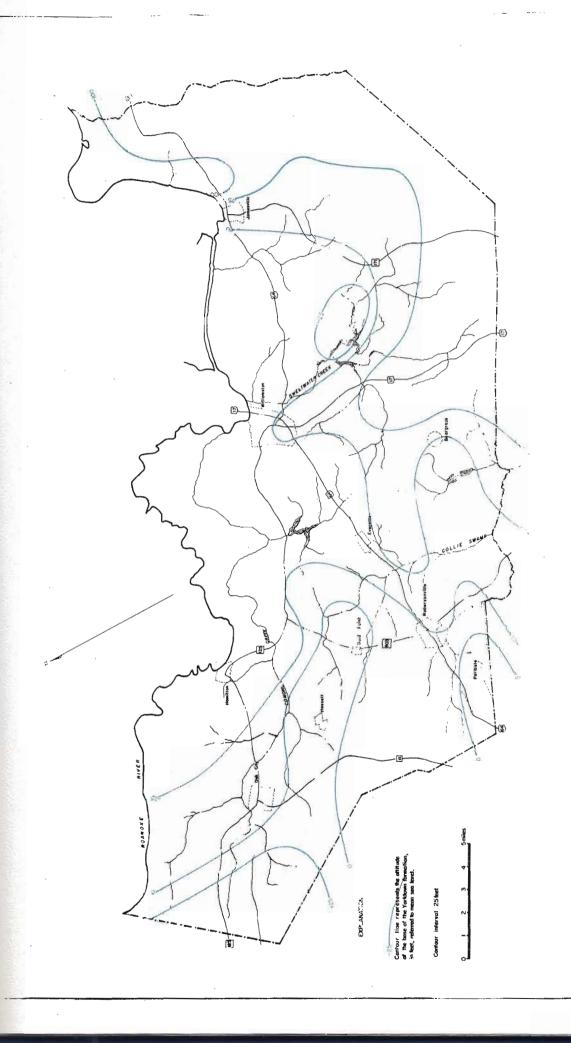
Upper Cretaceous Series

Tuscaloosa Formation. — The Tuscaloosa Formation unconformably overlies crystalline rock in most of Martin County. In the remainder of the county it overlies Lower Cretaceous sediments. The top of the formation strikes N. 51° E. and dips approximately 16 feet per mile to the southeast. The formation thickens toward the southeast, is approximately 200 feet thick in the western part of the county, and is approximately 300 feet thick in the eastern part.

Table 1. Depths Below Land Surface of Geologic Formations in Martin County

b indicates bottom depth of well-

		Miocene	20000	Pullincone		ipper dretacos		Tower	Crystelline
Mell Number	.colmornary	Yorktown Formation	Castle Layne Limestone	Pormation	Pactice Formation	Black Greek Formation	Pormogloss Pormogloss	Cretaceous	hauement Rock
71	Ü −20	20-112				112-390	390-614		6240
T 2	0-15	15-132	132-153	163-312	312-450	450-696	096-9755		
Al)-15	15-445							
A2	3-10	10-70				70-1035			
A4	5-10	10-855							
ئۂ	3-10	10-43		43-635				,	
20	J-15	15-80		80-67Ъ					
Α7	Ŭ - 5	5-83b							
AS	Ŭ -15	15-47				47 - 976			
A9	Ŭ -2 0	23-89		 	ช9ช	47-376			
AlO	J-24	24-51			51-77b				
All	0-20	20-75			1				
	1	1		(.6. 0.3	75-1035				
A12	0-25	25-80		80-91	91-1035				
A14	0-30	30-110		110-117;					
A15	J-20	20-90ъ							
A18	C-30	30-35Ь							
A19	0-20	20-90ъ							
A20	C−35	35-42b							
A21	0-30	30-35ъ							
A22	J-15	15-65	65-71	71-97Ь					
A-23	C-15	15-45		45~107b					
A24	ũ-20	20-65	65-75	75 - 976					
à25	0-20	20-22ь							
A26	2-5	5-70	70-96	96-1275					
ь 2 7	0-8	8-122	122-124b						
A2 8	5-10	10-110	110-115b						
A29	0-30	3C-95b							
A 3C	0-35	35-97	97~10Gb						
A31	0-15	15-118	118-1475						
A32	0-45	45-145b	110 1475	İ					
A23	0-15	15-96	96 - 975						
A34	Ŭ-20	26-55	30-375	'		Es OEb			
A35	6-20	20-74			7/ 1101	55-95b			
					74-1125				
A35	0-20	20-79			79-107ь				
1.37	0-20	20-30				80-107ъ			
7.33	0-305	10.15							
::1	0-40	40-60					60-345	24 5-335b	
21	0-50	50-155				155-370	370-415b		
DI	0-25	25-105		105-150	150-245	245-3205			
	0-65	45-120	120-160b						
20	0-70	70-210				110-2035			
24	0-70	70-105				105-1236			
127	0-31	31-,113				113-225	225-369 ხ		
1.50	0-25	35 - 125			125-2015				
:.60	0-40	40-120			130-250	250-290ნ			
10.7	0-30	30-125		125-150	150-275	275-3005			
2.	0-40	40-115		115-180	180-260	260-3205			
201	0-27	27-1.23		125-152	152-281	281-573	573 -7 025		
210	0-25	23-125		125-150	150-290	290-2905	5/5 /020		
21	0-20	20-120		129-160	150-250	29 0- 2605			
317	0-10	11 - 18	68-025		190-010	200-2500			
366	05	15-57	07-006						
406	0-10	40-120	120-160b						



MAP SHOWING CONTOURS ON THE BASE OF THE YORKTOWN FORMATION

The Tuscaloosa Formation consists of layers of white to red clay interbedded with layers of gray micaceous sand. Generally, the basal part of the formation is a massive clay bed about 75 feet thick. The center part consists of approximately 50 feet of interbedded sand and clay layers, and the upper part consists of 100 to 150 feet of clay overlain by 10 to 30 feet of sand (fig. 6).

Aquifer 1 is near the center of the Tuscaloosa Formation (fig. 5). Aquifer 2 occurs along the contact between the Tuscaloosa Formation and the Black Creek Formation.

In the western part of the county the chemical quality of water from these aquifers generally is acceptable for most uses. In the central and eastern parts of the county the water from Aquifers 1 and 2 contains high concentrations of chloride and, therefore, is too salty for most uses.

Black Creek Formation. -- The Black Creek Formation unconformably overlies the Tuscaloosa Formation in Martin County. The surface of the Black Creek Formation strikes N. 50° E. and dips about 14 feet per mile (figs. 5 and 6). The formation, which is approximately 325 feet thick throughout the central and eastern parts of the county, pinches out west of the Martin-Halifax County line (fig. 6). A projection of subsurface contacts indicates that the lower part of the Black Creek Formation crops out at Palmyra Landing, in Halifax County, in a cut along the south bank of the Roanoke River.

In Martin County the Black Creek Formation consists of interbedded layers of gray to brown clay and gray sand. The accessory minerals include: muscovite, pyrite, hematite, and glauconite. Lignitized wood fragments and shell fragments also are common in the samples.

A thick sand bed, designated Aquifer 3, occurs near the center of the Black Creek Formation. This aquifer is one of the most productive in the county. Another aquifer, designated Aquifer 4 occurs along the contact between the Black Creek Formation and the overlying Peedee Formation. Generally, water from both Aquifers 3 and 4 is of acceptable chemical quality for most domestic uses.

<u>Peedee Formation</u>. — The Peedee Formation in Martin County overlies the Black Creek Formation and underlies the Beaufort and Yorktown Formations (fig. 6). The surface of the Peedee Formation strikes N. 50° E.

250-266

Silt, gray, with sand, quartz, clear, very fine to fine-grained, rounded, 70 percent; sand, glauconitic, green, fine to medium-grained, rounded, 30 percent.

266-281

Sand, quartz, clear, very fine to finegrained, rounded, 70 percent; sand, glauconitic, green, fine to mediumgrained, rounded, 30 percent, with angular, medium-grained, brown quartz sand and few pyrite crystals.

281-296

Sand, quartz, clear, very fine to finegrained, rounded, 55 percent; sand, quartz, brown, medium-grained, angular, with few pyrite crystals, 15 percent; sand, glauconitic, green, fine- to medium-grained rounded, 30 percent.

296-301

Silt, gray, with rounded very fine to fine-grained clear quartz sand, 55 percent; sand, quartz, brown, medium-grained, angular, with few pyrite crystals, 15 percent; sand glauconitic, green, fine to medium-grained, rounded, 30 percent.

301-312

Sand, quartz, clear, very fine to finegrained, rounded, 55 percent; sand, quartz, brown, medium-grained, angular, with few pyrite crystals, 15 percent; sand, glauconitic, green, fine to medium-grained, rounded, 30 percent. Ostracoda identified from the 163-312 foot interval include:

Brachycythere Trachyleberis midwayensis (Alexander)
Trachyleberis prestwichiana (Jones and Sherborn)

Upper Cretaceous - Peedee Formation

312-317

Clay, black, with rounded, very fine to fine-grained clear quartz sand, 55 percent; sand, quartz, brown, medium-grained, angular, with few pyrite crystals, 15 percent; sand, glauconitic, green, fine to medium-grained, rounded, 30 percent.

	25
Depth (feet)	
317–327	Sand, quartz, clear, very fine to fine- grained, rounded, 55 percent; sand, quartz, brown, medium-grained, angular, with few pyrite crystals, 15 percent; sand glauconite, green, fine to medium- grained, rounded, 30 percent.
327 – 353	Clay, dark brown, with rounded, very fine to fine-grained, clear quartz sand, 55 percent; sand, quartz, brown, medium-grained, angular, with few pyrite crystals, 15 percent; sand, glauconitic, green, fine to medium-grained, rounded, 30 percent.
353–358	Sand, quartz, clear, very fine to fine-grained, rounded, 55 percent; sand, quartz, brown, medium-grained, with few pyrite crystals, 15 percent; sand, glauconitic, green, fine to medium-grained, rounded, 30 percent.
358 –3 68	Sand, quartz, white, very fine to coarse- grained, angular to well rounded, 90 percent; sand, glauconite, and phos- phatic, fine-grained, with muscovite flakes, 10 percent.
368-409	Clay, dark brown, with trace of angular to well rounded, very fine to coarsegrained white quartz sand, 90 percent; sand, glauconite, fine-grained, 10 percent; with trace of phosphatic sand, pyrite crystals, and muscovite flakes.
409-419	Sand, quartz, white, very fine to coarse- grained, angular to well rounded, 90 percent; sand, glauconite, fine-grained, 10 percent; with trace of phosphatic sand, pyrite crystals, and muscovite flakes.

419-424

Clay, brown, with trace of angular to well rounded, very fine to coarse-grained, white quartz sand, 90 percent; sand, glauconite, fine-grained, 10 percent; with trace of phosphatic sand, pyrite crystals, and muscovite flakes.

424-440

Sand, quartz, white, very fine to mediumcoarse-grained, angular to well rounded, 90 percent; sand, glauconite, finegrained, 10 percent; with trace of phosphatic sand, pyrite crystals, and muscovite flakes.

440-445

Sand, quartz, white, very fine to finegrained, subangular, 80 percent; sand, glauconite, green, fine-grained, 20 percent; trace muscovite. Ostracoda identified from the 312-445 foot interval include:

Brachycythere rhomboidalis (Berry)
Cytherella herricki Brown
Cytheridea (Haplocytheridea)

councilli Brown
Cythereis pidgeoni (Berry)
Eucythere curta Jennings
Velarocythere cacumenata Brown
Velarocythere eikonata Brown

<u>Upper Cretaceous - Black Creek Formation</u>

445-476

Clay, dark brown, with trace of subangular, very fine to fine-grained, white quartz sand, 80 percent; sand, glauconite, green, fine-grained, 20 percent; trace muscovite.

476-527

Clay, verigated, cream to brown, with trace of subangular, very fine to fine-grained white, quartz sand, 80 percent; sand, glauconite, green, fine-grained, 20 percent; trace muscovite with few shell fragments.

527-553

Sand, quartz, white, fine to medium-grained, angular to subrounded, 90 percent; sand, glauconite, green, fine-grained, 10 percent; trace marcasite crystals, muscovite flakes, chlorite, and lignitized wood.

553-568

Sand, quartz, white, medium to coarse-grained, angular to subangular; trace muscovite flakes, marcasite crystals, lignitized wood.

568-573

Clay, cream to tan.

Depth (feet)	
573–578	Sand, quartz, white, fine to coarse-grained, angular to subangular; trace muscovite flakes, marcasite crystals, lignitized wood.
578–583	Clay, tan
583–599	Sand, quartz, white, fine to coarse-grained, angular to subangular; trace muscovite flakes, marcasite crystals, lignitized wood.
599–604	Clay, brown.
604-614	Sand, quartz, white, very fine to medium- grained, subangular to subrounded; few muscovite flakes.
614–619	Clay, brown.
619-624	Sand, quartz, white, very fine to medium- grained, subangular to subrounded; few muscovite flakes.
624-629	Clay, brown.
629-640	Sand, quartz, white, very fine to medium- grained, subangular to subrounded; few muscovite flakes.
640–645	Clay, tan to brown.
645–660	Sand, quartz, white, very fine to medium- grained, subangular to subrounded; with few muscovite flakes and trace of shell fragments.
660–676	Sand, quartz, white, very fine to medium- grained, subangular to subrounded; with few muscovite flakes, trace of shell fragments, marcasite crystals, and 3 percent green, fine-grained glauconite.
676–686	Silt, gray, with trace of subangular to subrounded, very fine to medium-grained white quartz sand, few muscovite flakes, and marcasite crystals.

686-696

Sand, quartz, white, very fine to mediumgrained, subangular to subrounded, with few muscovite flakes, and marcasite crystals. Ostracoda identified from the 445-696 foot interval include:

Brachycythere Brachycythere nausiformis Swain
Brachycythere sphenoides (Reuss)
Cytheridea (Haplocytheridea)
monouthensis (Berry)
Cytheropteron (Eocytheropteron)
striatum Brown
Orthonotacythere sulcata Brown

Upper Cretaceous - Tuscaloosa Formation

696-706

Sand, quartz, white, very fine to mediumgrained, subangular to subrounded, with few muscovite flakes, and marcasite crystals.

706-711

Clay, red to gray, with trace of subangular, very fine to coarse-grained white quartz sand, abraded shell fragments, marcasite crystals, chlorite, and few quartz crystals ranging in color from white through yellow, green, brown, and smoky.

711-722

Sand, quartz, white, very fine to coarsegrained, subangular, with abraded shell fragments, marcasite crystals, chlorite, and a few quartz crystals ranging in color from white through yellow, green, brown, and smoky.

722-728

Clay, gray, and variegated red and white, with trace of subangular, very fine to very coarse-grained, white quartz sand, abraded shell fragments, marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky.

728-737

Sand, quartz, white, very fine to very coarse-grained, subangular with abraded shell fragments, marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky.

737-768

Sand, quartz, white, very fine to very coarse-grained, subangular, with abraded shell fragments, marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky, with lignitized wood fragments.

768-773

Clay, reddish brown, with trace of subangular, very fine to medium-grained,
white, quartz sand, abraded shell
fragments, marcasite crystals, chlorite, few quartz crystals ranging in
color from white through yellow,
green, brown, and smoky, with
lignitized wood fragments.

773-778

Sand, quartz, white, very fine to mediumgrained, subangular, with abraded shell fragments, marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky, with lignitized wood fragments.

778-783

Clay, reddish brown with trace of subangular, very fine to medium-grained, white, quartz sand, abraded shell fragments, marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky, with lignitized wood.

783-793

Sand, quartz, white, very fine to mediumgrained, subangular, with abraded shell fragments, marcasite crystal, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky, with lignitized wood.

793-799

Sand, quartz, white, very fine to mediumgrained, subangular with marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky, with lignitized wood.

799-855

Clay, reddish brown, with trace of subangular, very fine to medium-grained, white, quartz sand, marcasite crystals, chlorite, few quartz crystals ranging in color from white through yellow, green, brown, and smoky, with lignitized wood.

Depth (feet)	
855–860	Sand, quartz, white, very fine to fine- grained, subangular, with numerous hematite grains, few glauconite grains.
860–865	Sand, quartz, white, very fine to fine- grained, subangular, with numerous hematite grains, few glauconite grains, and shell fragments.
865-891	Clay, gray to brown, with trace of sub- angular, very fine to fine-grained, white, quartz sand, numerous hematite grains, few glauconite grains and shell fragments.
891–901	Sand, quartz, white to red, fine to very coarse-grained, angular to subrounded, with numerous medium to coarse-grained hematite grains, few abraded shell fragments.
901–911	Sand, quartz, white, very fine to fine- grained, subangular to subrounded, with few fine-grained hematite grains, few muscovite flakes.
911–916	Clay, red, with trace of sand, quartz, white, very fine to fine-grained, subangular to subrounded, with few fine-grained hematite grains, few muscovite flakes.
916–922	Sand, quartz, white, very fine to fine- grained, subangular to subrounded, with few fine-grained hematite grains, few muscovite flakes.
922–927	Clay, red, with trace of subangular to subrounded, very fine to fine-grained white, quartz sand, few fine-grained hematite grains, few muscovite flakes.
927–932	Sand, quartz, white, very fine to fine- grained, subangular to subrounded, with few fine-grained hematite grains, few muscovite flakes.
932-937	Clay, red, with trace of subangular to subrounded, very fine to fine-grained, white, quartz sand, few fine-grained hematite grains, few muscovite flakes.

937-942

Sand, quartz, white, very fine to fine-grained, subangular to subrounded, with few medium-grained quartz grains, few fine-grained hematite grains, few muscovite flakes.

942-957

Clay, cream to brown, with trace of subangular to subrounded, very fine to fine-grained white, quartz sand, few medium-grained, quartz grains, fine-grained hematite grains, and muscovite flakes.

957-963

Sand, quartz, white, very fine to coarsegrained, well sorted, subangular to subrounded, with few fine-grained hematite grains, few muscovite flakes.

963-973

Clay, cream to red, with trace of subangular to subrounded, very fine to coarse-grained, white quartz sand, few fine-grained hematite grains, few muscovite flakes.

973-983 (bottom) Sand, quartz, white, very fine to coarsegrained, subangular to subrounded, with few fine-grained hematite grains, few muscovite flakes, chlorite, and few white, yellow and pink clay fragments. No microfauna in 696-983 foot interval.

HYDROLOGY

Hydrologic Cycle

The supply of water on earth is kept in endless circulation between the land, seas, and atmosphere by energy supplied from the sun. This circulation has been termed the hydrologic cycle and the study of its various phases is the science of hydrology. In the hydrologic cycle, water is evaporated mostly from bodies of water and carried by the atmosphere until it condenses and falls as precipitation. Of the precipitation that falls over land surfaces, part is returned to the atmosphere by evaporation and the transpiration of plants, part drains overland into lakes and streams and returns to the oceans, and part seeps into the ground.

Ground Water

Some of the water which seeps into the ground is evaporated or transpired by plants and some percolates downward to the zone of saturation. Water in the zone of saturation, the zone in which all of the pore spaces are filled with water under positive hydrostatic head, is termed ground water.

In the zone of saturation, ground water moves under the influence of gravity toward general areas of discharge, such as stream channels and the ocean, and to points of discharge, such as springs or wells. Where ground water only partially fills a permeable formation, its surface, which is at atmospheric pressure, is free to rise and fall, and the water is said to be under non-artesian or water-table conditions. The water table is the surface of the non-artesian ground water. Where ground water completely fills a permeable bed that is underlain and overlain by relatively impermeable beds, its surface is not free to rise and fall and the water is said to be under artesian conditions. The piezometric surface is the surface which represents the height to which water will rise in wells tapping a given artesian aquifer. The term "artesian" is applied to water that is under sufficient pressure to rise above the top of the permeable formation containing it, although not necessarily above land surface. Thus, all flowing wells would be artesian but all non-flowing wells would not necessarily be non-artesian wells.

A unit in the zone of saturation that is permeable enough to transmit usable quantities of water to wells or springs is called an aquifer. Aquifer is a relative term used to denote that the lithic unit is more permeable than overlying or underlying units and that it will transmit sufficient quantities of water for specific purposes. In this report, aquifers are permeable zones which will furnish to wells, at a minimum, sufficient quantities of water for most domestic purposes. Aquifers are separated by beds of low permeability which are called confining beds or aquicludes. Areas in which aquifers are replenished are called recharge areas and areas in which water is lost from aquifers are called discharge areas.

Ground water moves from recharge areas to discharge areas. Water moves from a place where its surface is high toward a place where its surface is low. Thus, water would move in the non-artesian aquifer from a place where the water-table is high toward a place where it is relative-Ty lower and in an artesian aquifer water would move from a piezometric high toward a piezometric low. In addition to the movements of water within an aquifer, there would be an interchange of water from aquifer to aquifer if the intervening confining layers were permeable enough and if the head differences between aquifers were great enough. Aquifers, confining layers, recharge areas, and discharge areas are illustrated in figure 8. The figure also shows the directions in which ground water would flow in response to the head differences in aquifers and between aquifers. Wells Al, A2, and A3, on the left side of figure 8, are in a recharge area. The water table stands higher than the piezometric surface of the lower two artesian zones so the movement of water would be downward, through the confining layers. Wells Cl, C2, and C3 are in a discharge area. Here the water table is lower than the piezometric surfaces and the water would tend to move upward from the artesian zones into the water table aquifer and seep into the stream. In this illustration the flow of water would be along the lines of the dashed arrows. This is a highly generalized section of the conditions found in Martin County where there are numerous recharge and discharge areas and several major artesian aquifers.

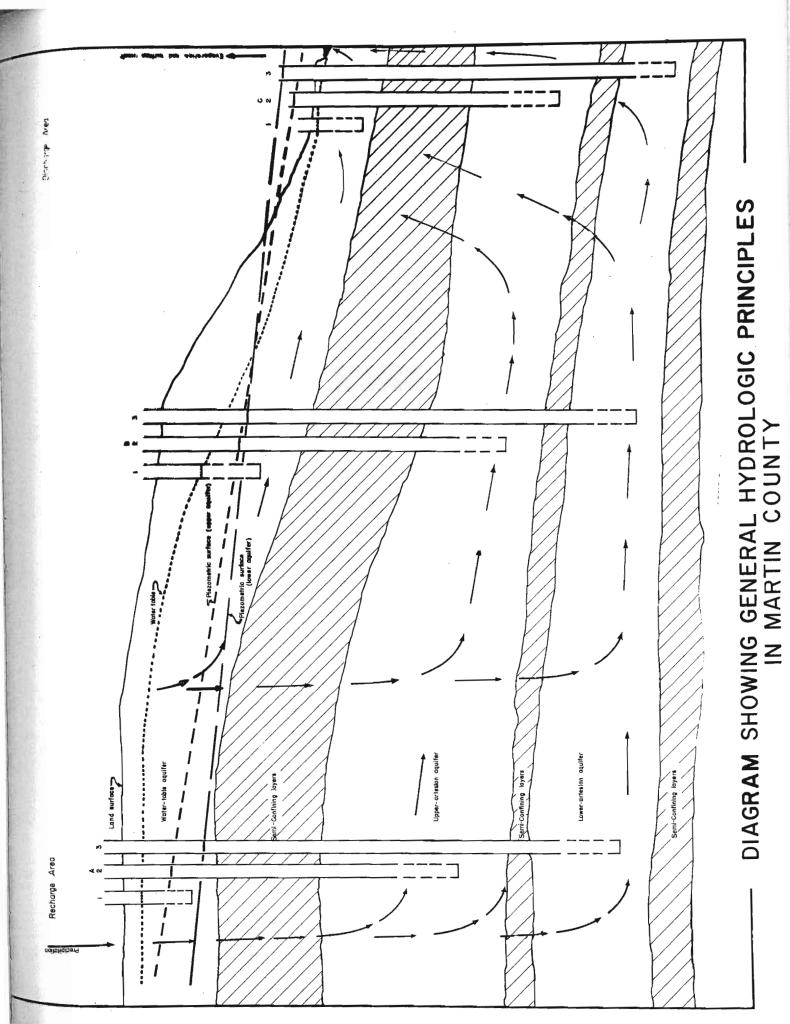
Artesian aquifers

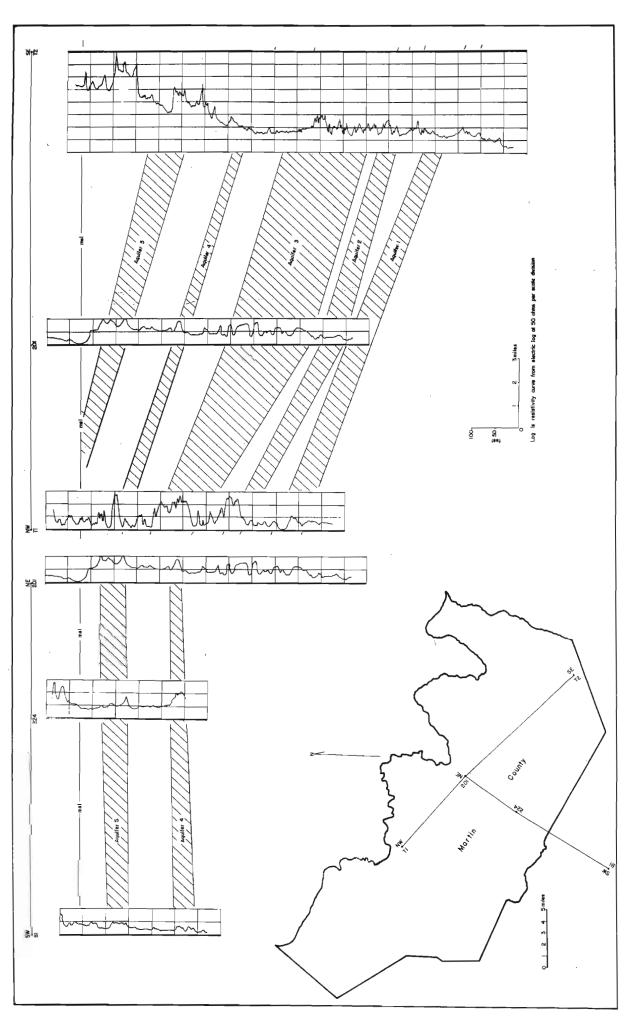
There are seven principal aquifers in Martin County, numbered 1 through 7. Aquifers 1 through 6 are artesian. The aquifers are numbered for ease of reference because some aquifers include parts of more than one geologic formation, and some formations include more than one aquifer. The aquifers are numbered consecutively from basement rock upward. Thus, Aquifer 1 is the first aquifer above basement rock and Aquifer 7 the topmost aquifer in the county. The aquifers are described in the following sections and are shown diagrammatically and in section on figures 5 and 9. The artesian aquifers, except Aquifer 1, pinch out along the highly permeable basal Yorktown unit. Therefore, water percolating through the Yorktown clay layers may readily move toward and into the upper aquifers. Because of this interconnection of aquifers, one piezometric map is satisfactory in depicting artesian head for the zone between depths of about 100 to 300 feet below the surface.

Aquifer 1

Aquifer 1 occurs near the middle of the Tuscaloosa Formation. It strikes N. 51° E. and dips to the southeast at approximately 16 feet per mile. The thickness of Aquifer 1 ranges from about 25 feet in the eastern part of the county to about 30 feet in the western part of the county. Its top is about 400 feet below land surface in the western part of the county and about 850 feet below land surface in the eastern part.

Aquifer 1 is composed of two sand beds and an intervening clay layer that thickens and thins across the county and may not be continuous throughout the county. The upper sand bed contains white, very fine to fine, subangular, poorly sorted quartz sand. The lower sand bed is composed of white, very fine to fine, subangular to subrounded, poorly sorted quartz sand. The optimum screen openings for naturally developed screened wells tapping Aquifer 1 is from 7/1000 to 10/1000 of an inch (slot size 7-10). The specific capacity of properly constructed, naturally developed (See p. 44), 2-inch diameter wells tapping Aquifer 1 is calculated as 0.6 gallons per foot of drawdown at the end of one day of pumping. Throughout most of the county, the water from Aquifer 1 contains high concentrations of chloride which makes it undesirable for most uses.





CORRELATION OF AQUIFERS BY ELECTRIC LOGS

Aquifer 2

Aquifer 2 includes the base of the Black Creek Formation in the western part of the county, and the upper part of the Tuscaloosa Formation in the eastern part.

Aquifer 2 strikes N. 49° E. and dips at about 14 feet per mile to the southeast. Aquifer 2 is about 60 feet thick in the western part of the county and about 45 feet thick in the eastern part. Its top is about 350 feet below land surface in the western part of the county and about 675 feet below land surface in the eastern part.

Aquifer 2, a sand bed containing a few thin clay stringers, is composed of white, very fine to medium, subrounded to subangular, poorly sorted quartz sand. The sand includes a small percentage of muscovite flakes and fine-grained marcasite crystals. The optimum screen openings for naturally developed screened wells tapping Aquifer 2 should be about 12/1000 inch (slot size 12). The specific capacity of properly-constructed, naturally developed, 2-inch diameter wells tapping Aquifer 2 is calculated as 0.7 gallons per foot of drawdown at the end of one day of pumping. Aquifer 2 is tapped by few wells because the water contains concentrations of chloride in excess of 400 parts per million throughout most of the county.

Aquifer 3

Aquifer 3 is perhaps the most productive and widely used aquifer in Martin County. Figure 5 shows that this aquifer in the central and western sections of the county is in part of the Black Creek Formation, and in the eastern section, is in the upper part of the Tuscaloosa Formation.

Aquifer 3 strikes N. 50° E. and dips toward the southeast at about 14 feet per mile. Aquifer 3 is about 80 feet thick in the western and central parts of the county. From the central part to the eastern part of the county, this zone thickens and reaches a thickness of about 150 feet along the eastern edge of the county. A map, figure 10, has been prepared to show the depths, below land surface, of the tops and bottoms of the principal aquifers. The top of Aquifer 3 crops out along the Roanoke River in the northwestern part of the county and is about 550 feet below land surface in the southeastern part.

Aquifer 3 is composed of thick sand beds containing numerous thin clay layers. The sand in the top 50 feet of the aquifer is white, medium to coarse grained, angular to subangular, poorly sorted, quartz sand containing a trace of muscovite flakes, marcasite crystals and lignitized wood fragments. The lower part of the aquifer is composed of the same material except that the sand grains are smaller and range from medium to very fine grained. The optimum screen openings for naturally developed screened wells tapping the upper part of Aquifer 3 should be 20/1000 inch (slot size 20). The screen openings for the lower part of Aquifer 3 should be about 12/1000 inch (slot size 12). The specific capacity of a properly-constructed, naturally developed, 2-inch diameter well tapping Aquifer 3 is calculated as 6.0 gallons per foot of drawdown at the end of one day of pumping.

The chemical quality of water from Aquifer 3 generally is acceptable for most uses in the western and central parts of the county. In the eastern part of the county, water from this aquifer usually contains concentrations of chloride in excess of 400 parts per million and is not desirable for most uses.

Aquifer 4

Aquifer 4 occurs along the Black Creek-Peedee Formation contact. In the eastern part of the county, the upper part of Aquifer 4 is in the base of the Peedee Formation but in the remainder of the county Aquifer 4 is in the Black Creek Formation. The strike averages about N. 50° E. and dip of Aquifer 4 to the southeast averages about 14 feet per mile. Aquifer 4 ranges in thickness from about 40 feet in the western part of the county to about 60 feet in the eastern part. The top of Aquifer 4 pinches out along the base of the Yorktown Formation, just west of Oak City, in the western part of the county. In the eastern part, the top of the aquifer is about 400 feet below land surface. The depths of the top and bottom of Aquifer 4, below land surface, are given in figure 10.

The aquifer is composed of a thick sand bed, overlain and underlain by massive clay beds. The sand in this aquifer is white, very fine to coarse, angular to subangular, poorly sorted quartz sand containing a trace of phosphatic and glauconitic sand, pyrite crystals

DEPTH OF PRINCIPAL AQUIFERS IN MARTIN COUNTY

and muscovite flakes. The optimum screen openings for naturally developed, screened wells tapping this aquifer is about 16/1000 inch (slot size 16). The specific capacity of properly-constructed, naturally developed, 2-inch diameter wells tapping Aquifer 4 is calculated as 1.5 gallons per foot of drawdown after one day of pumping. The chemical quality of water from Aquifer 4 generally is acceptable for most uses.

Aquifer 5

Aquifer 5 occurs in and includes most of the Beaufort Formation. The aquifer ranges in thickness from less than a foot along the basal pinchout to approximately 100 feet in the eastern part of the county. The aquifer strikes approximately N. 40° E. and dips at about 20 feet per mile to the southeast. The base of the aquifer pinches out along an undulating, north-south line which passes through Everetts, in the central part of the county. The top of the aquifer pinches out along a line, roughly parallel to the strike, which passes approximately half way between Williamston and Jamesville. The depth of the top and bottom of Aquifer 5 below land surface is shown in figure 10.

Aquifer 5 is composed of about 60 percent, silt-sized to medium-grained, green to black, glauconite sand and about 40 percent, silt-size to fine-grained, clear quartz sand. The sands are poorly sorted. The proper screen openings for naturally developed, screened wells tapping Aquifer 5 is about 8/1000 inch (slot size 8). The specific capacity of a properly constructed, naturally developed, 2-inch diameter well tapping Aquifer 5 at its greatest thickness, is calculated as 0.4 gallons per foot of drawdown after one day of pumping. The chemical quality of the water generally is acceptable for most uses.

Aquifer 6

Aquifer 6 occurs in and includes most of the Castle Hayne Limestone, and, where the Castle Hayne is overlain by the Yorktown Formation, it also includes a basal shell bed in the Yorktown Formation. The base of Aquifer 6 pinches out along an undulating, north-south line which passes just east of Williamston. The thickness of the aquifer ranges from less than a foot along the pinchout to about 40 feet along the eastern edge of the county. The top of Aquifer 6 strikes approximately N. 45° E. and dips toward the

southeast at about 18 feet per mile. The depths of the top and bottom of Aquifer 6 below land surface is shown in figure 10.

Aquifer 6 is composed of shell and shell limestone beds. The limestone is sufficiently indurated so that open-end wells may be constructed. The specific capacity of open-end, 2-inch diameter wells, where Aquifer 6 is 10 feet thick, is calculated as 3.0 gallons per foot of drawdown after one day of pumping. The water from Aquifer 6 generally is hard to very hard, otherwise it is of acceptable chemical quality for most uses. Aquifer 6 is the uppermost of the artesian aquifers in Martin County. Overlying these artesian aquifers are confining layers consisting of massive clay beds of the Yorktown Formation. The Yorktown Formation is overlain by sediments of post-Miocene age that are the non-artesian aquifer.

Non-artesian (water-table) Aquifer

Aquifer 7

Aquifer 7, the non-artesian aquifer, is composed of sand, silt, and clay of post-Miocene age. The aquifer extends from land surface to depths ranging from 5 to 70 feet in the county. Sand and silt beds predominate, but there are numerous clay lenses and clay is disseminated through the sand beds.

The sand beds in Aquifer 7 are composed of very fine to coarse-grained, well-rounded to angular, quartz sand that is frequently stained by iron oxides. The sand beds are heterogeneous and lack continuity; therefore, it is not practical to determine a standard screen size for wells tapping this aquifer. The optimum screen size should be determined, individually, for each well constructed.

Many 2-inch wells have a specific capacity of less than 0.1 gallon per foot of drawdown and none had a specific capacity of more than 0.3 gallon per foot of drawdown after one day of pumping. Therefore, the use of Aquifer 7 is limited to small diameter wells having small yields or large diameter wells that have a large reservoir storage capacity. Most farm ponds in the area that have no surface inflow or outflow, are essentially large diameter wells developed in Aquifer 7.

Aquifer 7 contains water having two objectionable characteristics. The water may have high concentrations of dissolved iron. This results in the staining of laundry, plumbing fixtures, and cooking utensils. The second objectionable characteristic is that water from Aquifer 7 is easily polluted by organic and inorganic sources such as septic tanks or chemical fertilizers.

Quantitative Studies

The quantitative studies of ground water in Martin County were divided into two major phases. The first phase, which included pumping tests to determine the water transmitting and water storage characteristics of various aquifers, was designed to determine the availability of ground water and the best method for developing water supplies. The second phase, using data obtained from the first phase and piezometric maps, was designed to calculate the areas and amounts of recharge and discharge within the county. The result is a calculation of how much ground water is available to both the individual supply well and to the area as a whole.

Water Storage and Water Transmitting Capacities of Aquifers

The water-transmitting capacity (coefficient of transmissibility) and the water-storage capacity (coefficient of storage) for the principal aquifers in Martin County were calculated from pumping and recovery tests. The coefficient of transmissibility is the quantity of water, in gallons per day, that will move through a vertical section of the aquifer one foot wide and extending the full saturated thickness under a hydrologic gradient of one foot per foot at the prevailing temperature of the water. The coefficient of storage is defined by the Geological Survey as the volume of water the aquifer releases from or takes into storage per unit surface area of the aquifer per unit change in the component of head normal to that surface. Determination of the transmissibility and storage coefficients for an aquifer may be used to calculate the water yielding capacity of individual wells tapping the aquifer for any reasonable period of pumping and the effect of drawdown in relation to distance from the pumped well.

The withdrawal of water from an aquifer causes water levels to decline in the vicinity of the pumped well. As a result of this

decline, the water table or piezometric surface assumes the approximate shape of an inverted cone having its apex at the center of withdrawal (fig. 11). The shape, size, and rate of growth of this "cone of depression" depends upon several factors. Among these are: (1) the transmissibility and storage coefficients of the aquifer; (2) the rate of pumping; (3) the duration of pumping; and (4) the natural discharge salvaged by increased recharge resulting from the decline in water levels near recharge areas. The vertical distance that water levels are lowered at any point by pumping is the "drawdown". The specific capacity of a well is the amount of water in gallons per minute that the well will yield for each foot of drawdown in the well for a specified period of pumping. As amount of drawdown is approximately proportional to the pumping rate, a well may be evaluated according to its specific capacity. The theoretical specific capacity of a well is the amount of water in gallons per minute that the well should yield for each foot of drawdown within the well, for a specified period of pumping if there are no well losses (B fig. 11). The specific capacity of a well is usually somewhat less than the theoretical specific capacity of the well due to head losses caused by turbulent flow through the well screens or poor well development or construction (C fig. 11).

Aquifer Studies

Pumping tests were conducted using wells tapping each aquifer in Martin County. The pumping tests ranged from elaborate, long-term tests to simple measurements of yield and drawdown during the construction of domestic wells. Table 2 gives the average hydrologic properties obtained from the tests, for each aquifer in the county except Aquifer 7, the water-table aquifer, for which the maximum rather than average properties are given. The table lists values for the theoretical specific capacities for 2-inch diameter wells and 36-inch gravel pack wells, assuming no well losses. These diameters were selected because most domestic wells are between $1\frac{1}{4}$ -inches and 3-inches in diameter and most municipal or industrial wells are gravel packed to approximately 36 inches.

The theoretical specific capacities of wells tapping more than one zone may be determined by adding the capacities of the zones tapped. For example, a 2-inch diameter well tapping Aquifers 5 and 6 should have a theoretical specific capacity of 3.4 gallons per foot of drawdown (table 2).

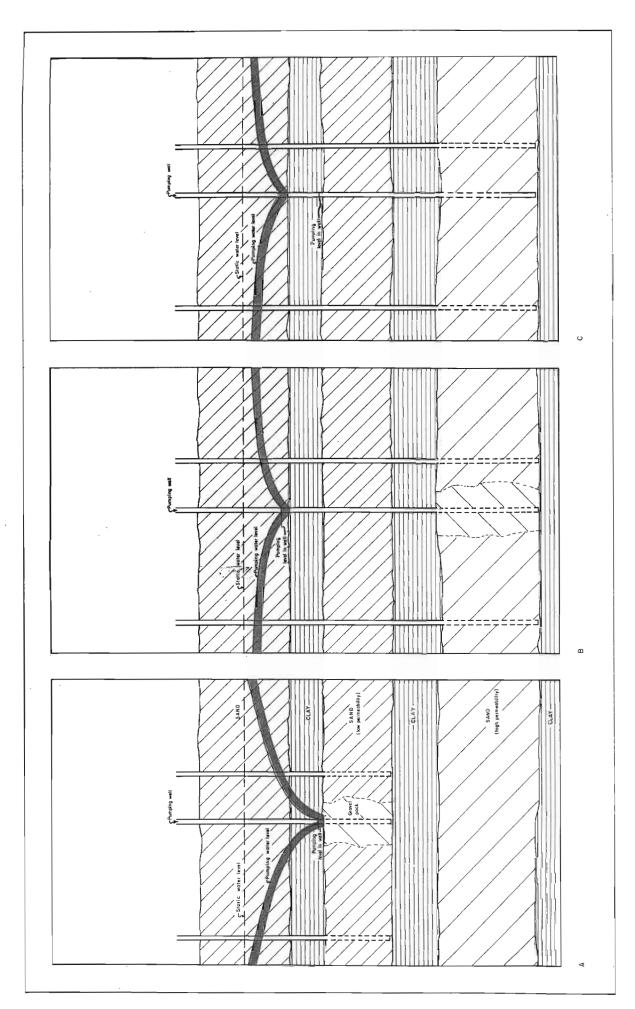


DIAGRAM SHOWING CONE OF DEPRESSION FORMED DURING PUMPING

TABLE 2. Average Quantitive Values for Aquifers in Martin County (Assuming no well losses)

Aquifer	Т	Aquifer Thickness (feet)	Pf		etical Capacity 36" diameter well
1	1,400	30	47	0.6	1.1
2	1,600	50	32	0.7	1.2
3	18,000	105	171	6.0	10
4	4,000	42	95	1.5	1.9
5	1,200	40	30	0.4	0.6
6	10,000	.9	1,100	3.0	4.0
7 *	700	26	27	0.3	0.5

^{*} Maximum measurement

T - Transmissibility

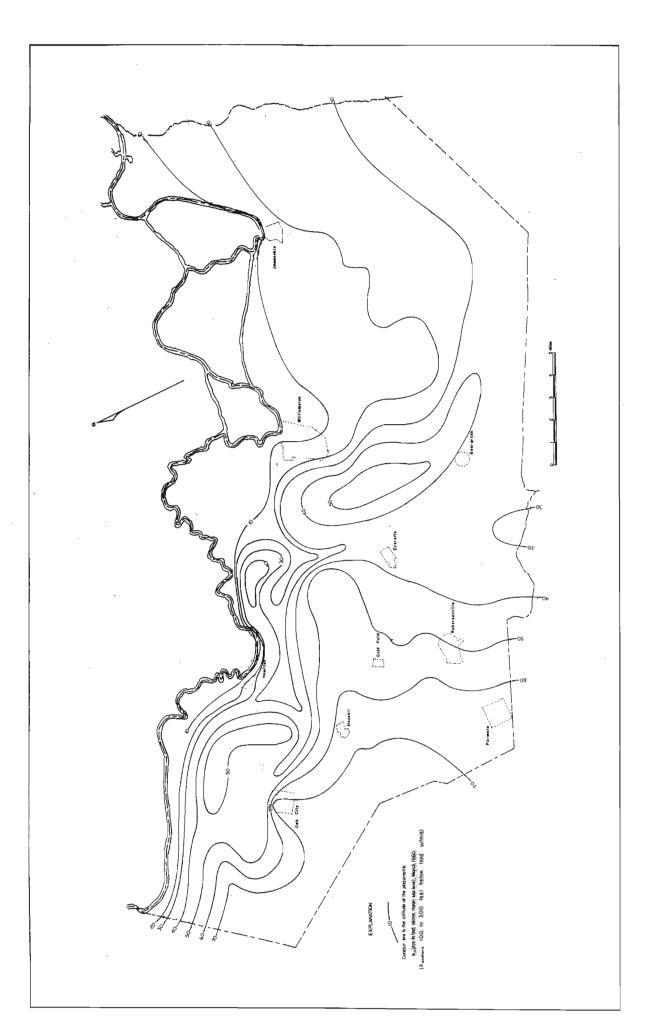
 $^{{\}tt PF - Field \ Permeability}$

The table may be used to compare the head losses of a well already constructed to a theoretically perfect well or it may be used to determine what zones to tap in constructing a new well. One may compare the losses of a well already constructed with a theoretically perfectly constructed well by measuring the drawdown and discharge of the well for a few hours to one day and calculating the specific capacity of the well. The specific capacity may be determined by dividing the discharge, in gallons per minute, of the well by the drawdown, in feet, in the well. If the specific capacity is much lower, due to losses, than the theoretical specific capacity listed for an aquifer, frequently, additional well development by the driller will considerably improve the yield of the well.

The quantitative studies in Martin County were designed to determine the average daily recharge to and discharge from the artesian aquifers between 100 feet and 300 feet below land surface in the county. The recharge or discharge for an area may be calculated by determining the size, shape, and transmissibility of the aquifers and the differences in head within the aquifers. In Martin County the size and shape of the aquifers were determined by geologic studies, the transmissibility was determined for the aquifers by conducting pumping tests and the head differences were determined by constructing a piezometric map of the county during a period when water levels were approximately average.

The piezometric map (fig. 12) was constructed by measuring and referring the water levels in wells to a common datum plane, mean sea level. As may be seen in figure 12, there are areas in the county where high-level contours close partially in Martin County and partially in adjacent counties. These areas of high piezometric surface are recharge areas where water is percolating downward into the artesian aquifers. The areas of low piezometric surface are discharge areas where water is leaking upward into swamps and streams. In the aquifers water flows from piezometric highs to piezometric lows as a result of the force of gravity.

. The average daily recharge is calculated by determining the amount of water moving away from recharge areas within the county. The average daily recharge in Martin County was calculated on the basis of



PIEZOMETRIC MAP OF MARTIN COUNTY, MARCH 1960

differences of piezometric head shown on figure 12 and the distribution of areas of recharge shown on figure 13. The figure calculated for recharge to aquifers between 100 feet and 300 feet below land surface in Martin County is 22 million gallons per day (mgd). This figure represents only recharge from precipitation and probably is conservative at that. It does not include the volume of water in the overlying nonartesian aquifer which could be induced to move downward by additional development of the artesian aquifers.

The average daily discharge for the same aquifers, calculated by determining the amount of water moving into discharge areas in the county is 21 mgd. The figures for recharge and discharge should be approximately equal if there is no outflow through the aquifer to other areas. From the piezometric map there appears to be one area near Parmele where there is outflow through the aquifer into Pitt County. The outflow and water use probably accounts for the discharge rate calculated being slightly less than the recharge rate calculated.

Comparison of average daily recharge to the aquifers between 100 and 300 feet below land surface datum and the volume of precipitation on the recharge areas (fig. 13) indicates that at least one quarter of the precipitation on the recharge areas is recharged to the 100 to 300 foot zone. In areas of discharge and transmission (fig. 13) little or no precipitation reaches the artesian aquifers; rather, any precipitation that reaches the zone of saturation moves through the water-table aquifer to points of discharge.

The protection of recharge areas is vital to the development of ground-water supplies in the county. Extensive drainage in these areas would affect substantially the total amount of ground water available for development. Also, long-lived contaminants should not be introduced into the recharge areas. For example, if soluble contaminants were introduced in area "E" (fig. 13), they could reach the Williamston municipal well field, located about 3 miles northeast of area "E", in approximately 80 years under present hydrologic conditions.

The drainage of areas between the recharge and discharge areas would probably result in the salvage of some of the runoff in recharge areas. The drainage of discharge areas would increase the flow from recharge to discharge areas by increasing the head differences, and,

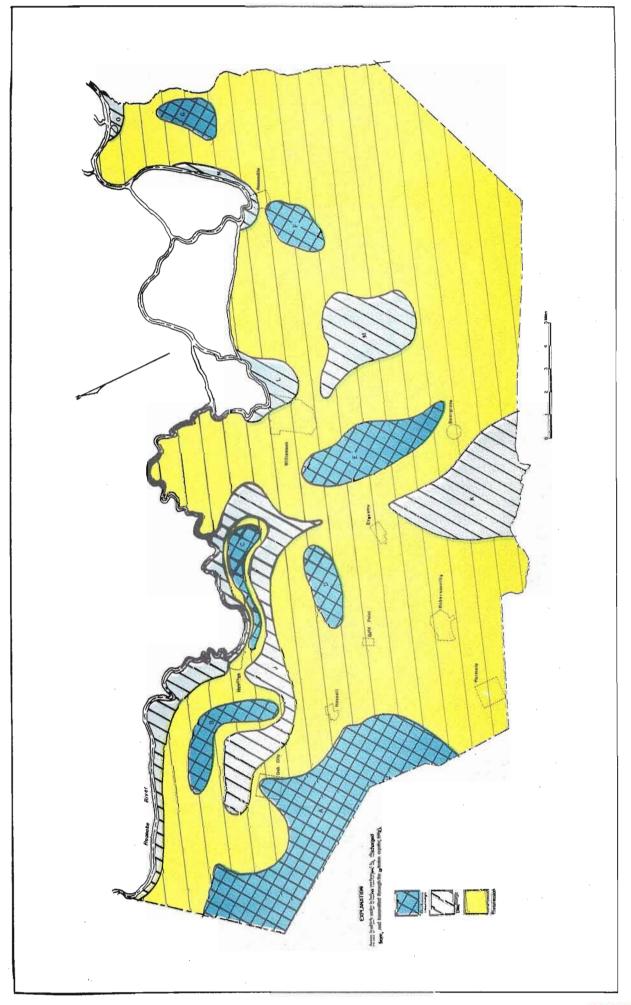
very likely, drainage would be difficult if not nearly impossible. For example, Sweet Water Swamp, Area "M" in figure 13, could be drained only with the greatest difficulty. The drainage of this area would require a system of ditches capable of handling about 6 mgd of base flow in addition to the total average daily rainfall for that area.

Generally, in Martin County it is more desirable to develop wells or well fields in the discharge areas. By developing in discharge areas a part of the normal leakage from the aquifer is utilized. Also, ground water in the discharge areas is usually of more desirable chemical quality than the water from recharge areas; this apparently anomalous relationship is discussed in sections on "Iron" and "Hardness".

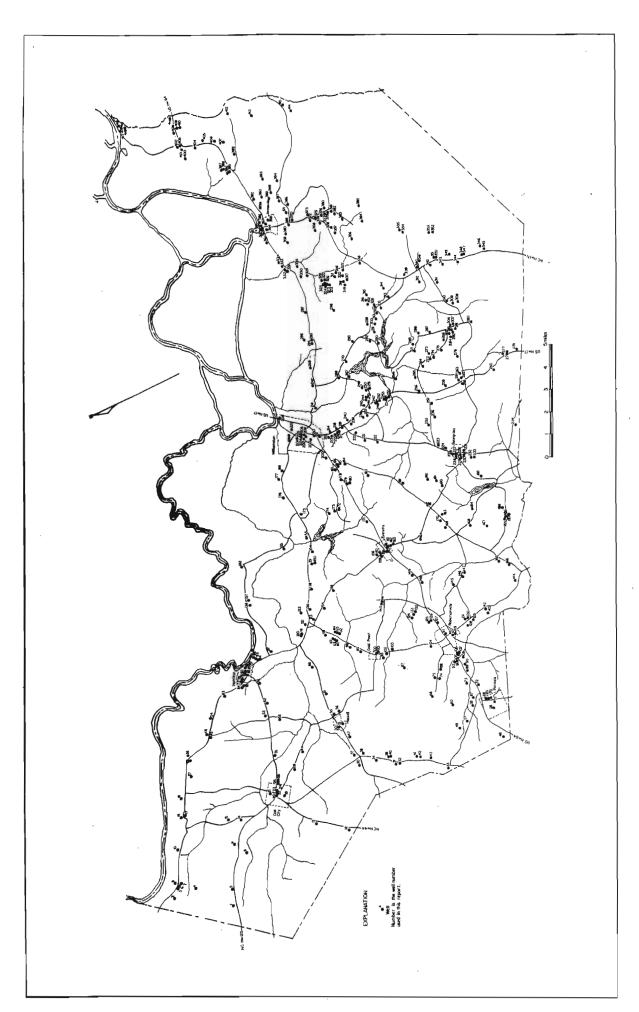
Wells

Domestic, municipal, and industrial wells were the source of most of the hydrologic and geochemical data collected in Martin County. Approximately 420 wells were inventoried. They are located in figure 14. The data from these wells are given in tables 3 and 4. Table 3 contains the physical data and table 4 contains the results of chemical analysis of water samples from the wells.

Generally, the wells are numbered from west to east across the county for the convenience of the reader in locating them on the map (fig. 14). Approximately 70 percent of the artesian wells in the county were inventoried, and the inventory network was designed to give comprehensive coverage, both areal, and with depth. A study of the data from the well inventory indicates some of the reasons why problems are encountered in obtaining wells capable of yielding adequate quantities of water of good chemical quality. Frequently, wells have a low specific capacity because of inadequate construction or development methods even though they tap very productive aquifers. Also, wells may produce water of undesirable chemical quality because they tap the least desirable of several aquifers in a given area. The following section on well construction is included to help minimize some of these problems.



AREAS OF RECHARGE, DISCHARGE, AND TRANSMISSION IN THE ARTESIAN AQUIFERS



Well construction

The types of wells constructed are governed by the relationship between the requirements of well owner and the cost of the well. The requirements of the owner include both quantity and quality of water required. Generally, there is no place in Martin County where an adequate well, yielding water of desirable quality, for most purposes cannot be developed. Since the individual must determine what his requirements are, the following section will describe the various methods of well construction along with advantages and disadvantages of each. For the purpose of this discussion, the following classifications are used: Cost: low - \$10.00 to \$250.00; medium - \$250.00 to \$1,000.00; high - more than \$1,000.00. Yield: low - 1 to 10 gpm; medium - 10 to 40 gpm; high - more than 40 gpm. Quality: good - does not require treatment for most domestic use; poor - requires some type of treatment for most domestic uses.

The types of wells generally constructed in Martin County are shown in figure 15. These wells fall into three general categories; dug or bored wells, driven wells, and drilled wells. In figure 15 the well on the left is a dug or bored well, the next well is a driven well, and the remainder of the wells in the illustration are drilled wells.

Dug or bored wells are constructed by digging or boring a large diameter hole, casing the hole with lengths of concrete or tile casing, grouting around the outside of the casing, and installing a pump. These wells range in depth from 10 to about 50 feet in Martin County. Usually they are 24, 30, or 36 inches in diameter. The price of these is low, the yield is generally low, and the quality of water is usually poor. Dug wells are usually constructed where high yield and good quality of water may be sacrificed for economy.

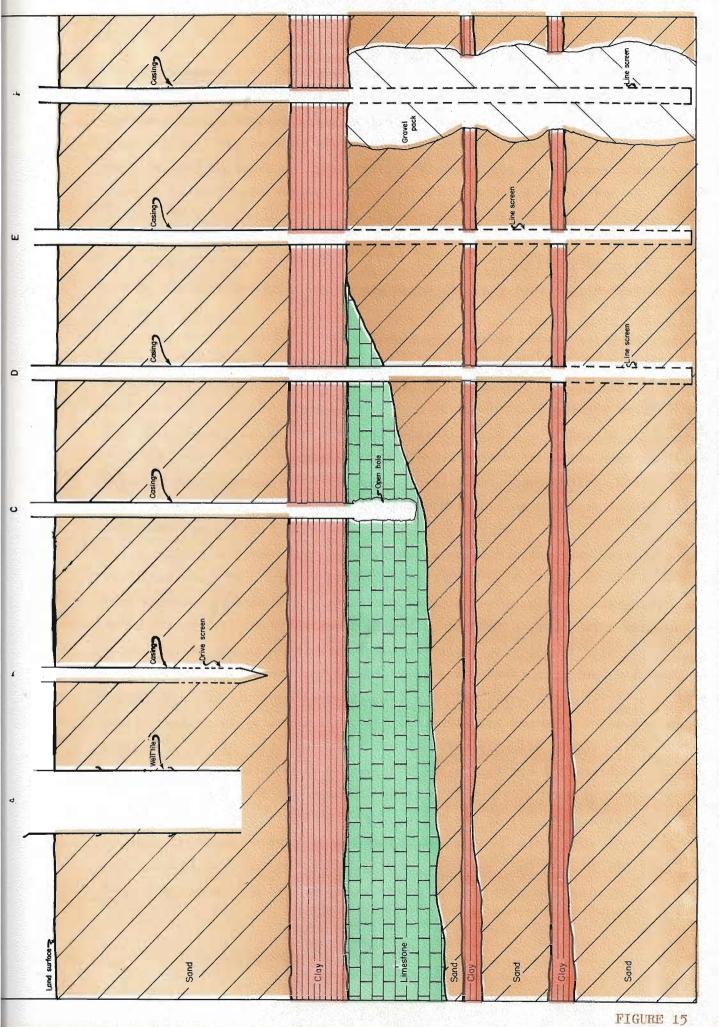
Driven wells are constructed by equipping a length of steel well casing with a special pointed drive screen which is driven down into a sandy water-bearing zone, surged to clear the fine sandy material from the screen, and then equipped with a pump. Driven wells in Martin County are usually l_4^1 to 2 inches in diameter and range from 10 to 50 feet in depth. These are low cost wells, usually having a low yield and poor quality of water. Driven wells, like dug wells, are usually

constructed where high yield and good quality of water may be sacrificed for economy.

Drilled wells are constructed by several methods in Martin County. Among these methods are: 1. the open-end method, 2. the single-screen natural-development method, 3. the multiple-screen natural-development method, and 4. the multiple-screen gravel-pack method. The type of well resulting from these methods of construction are shown in figure 15. Single-screen wells may also be gravel packed but none of this type were inventoried.

Well "C" in figure 15 illustrates the open-end type of well construction used in Martin County. In this method of construction, a hole is drilled from land surface into consolidated rock, such as limestone, a steel casing is set to the top of the consolidated rock and grouted in place, and the well is cleaned by pumping. This type of construction is restricted to areas which are underlain by sediments consolidated sufficiently that the hole will remain open below the bottom of the casing. Only the eastern one-third of Martin County is underlain by this type of formation. Here the Castle Hayne Limestone (Aquifer 6) underlies the area and numerous open-end wells tap this formation. The open-end wells are medium cost wells, usually of medium to high yield and the water is generally of poor quality. Water from open-end wells in the county is generally hard to very hard. The open-end wells are usually constructed in the eastern part of the county where medium to high yield is required and where hardness of water is relatively unimportant or is remedied by use of water softeners.

The single-screen naturally developed well is illustrated by Well "D" in figure 15. In this method of construction a hole is drilled from land surface down through a sandy aquifer, a steel casing with a screen on the lower end is set in the hole so that the screen is in the sandy aquifer, and the well is developed to wash fine sand and clay out of the aquifer around the screen. In this method of construction it is important that the screen openings are large enough to allow about 50 percent to 80 percent of the fine sand in the aquifer to wash into the well. By developing the well with pumping and surging, the fine material of the aquifer is washed out of the well and formation,



TYPES OF WELLS CONSTRUCTED IN MARTIN COUNTY

leaving coarse, more permeable material packed around the screen. This type of construction should produce a well of medium cost, usually of medium yield, producing water of good chemical quality, if properly constructed. The common problems in the construction of this type of well usually include: 1. poor well development due to improper screen size or insufficient surging and pumping, 2. setting the screen so that part of the screen openings are blocked by clay layers, 3. screen set in aquifer containing water of poor chemical quality.

Most of the single-screen, naturally-developed wells in Martin County are 2 inches in diameter and range from 125 feet to 350 feet in depth.

Multiple-screen naturally developed wells are constructed in the same manner as single-screen naturally developed wells except that screens are placed in the casing to tap more than one aquifer (see Well "E", figure 15). This type of construction should produce wells of medium cost, high yield and good chemical quality of water in Martin County. The problems of this type of construction are same as those of constructing single-screen naturally developed wells. Wells of this type in Martin County are usually 2 inches to 6 inches in diameter and range from 150 feet to 450 feet in depth.

Multiple-screen gravel-packed wells are illustrated by Well "F" in figure 15. This type of well is constructed by drilling a hole from land surface through several aquifers, reaming or washing a large diameter hole in each aquifer to be screened, setting a casing with screens spaced in the casing at each reamed aquifer, pumping gravel into the large diameter holes outside of the screens and developing the well to remove drilling clay and fine material from the graveled zones. The gravel packing effectively increases the radius of the well in each aquifer which, in turn, increases the specific capacity of the well. The cost of constructing this type of well is high but the yield of the wells is high and the chemical quality of water is usually good. Generally, only high yield municipal or industrial wells are constructed in this manner because of the cost. In Martin County, this type of well is usually 6 inches to 12 inches in diameter, 300 feet to 450 feet deep, and yields 400 to 600 gpm of water of good chemical quality.

GEOCHEMISTRY

All of the ground water in Martin County contains some dissolved matter which imparts certain chemical characteristics to the water. The amount and type of dissolved matter in a water sample is determined by chemical analysis of the water. The information obtained from approximately 400 chemical analyses of water, from the different aquifers, at various places within an area, may be related to the chemical composition of the aquifers and to the rate and direction of movement of water through the aquifers. The study which relates water chemistry to geology and hydrology is termed geochemistry.

A study of geochemistry in Martin County is important because it allows the development of aquifers in which water has the most desirable chemical characteristics and because it helps in tracing the movement of ground water within the county.

Quality of Water

Rainfall over the recharge areas of Martin County is relatively free of dissolved solids but does contain absorbed gasses, such as carbon dioxide and oxygen, which makes the water slightly acid or corrosive. The slightly acid water is capable of dissolving certain solids from the host rock through which it flows.

The salts of the common metals make up a large percentage of the dissolved solids in ground water. The chemical analysis of a water sample measures the individual constituents present in the solution. The proportionate amounts and chemical relationships of the ions in the solution determines such reported factors as hardness, specific conductance, and hydrogen-ion concentration. The following discussion concerns the more important chemical constituents in ground water in Martin County, as normally reported in water analyses.

Iron (Fe)

Iron compounds are common in the surficial sands in the county and iron is readily dissolved by the slightly acid water in the water table. Water containing less than 0.3 ppm of iron is generally acceptable for most uses. Concentrations of iron in excess of 0.3 ppm will usually stain laundry, food, utensils, and plumbing fixtures.

The iron stain, a reddish-brown deposit, forms when the water containing iron salt is exposed to the air. The dissolved iron may usually be removed from domestic supplies by using commercial type filters in the pumping system.

In Martin County the water-table aquifer yields water with objectionable amounts of dissolved iron. Almost all water-table wells yield water with iron concentrations ranging from 0.5 to 10 ppm. Therefore, this is the greatest objection to using dug or driven water-table wells.

The concentration of dissolved iron in water from aquifers between 100 feet and 300 feet below land surface is shown in figure 16. The water having greatest concentrations of dissolved iron occurs immediately under recharge areas (fig. 13). This is because water from the water-table aquifer, containing high concentrations of dissolved iron, is moving downward in these areas to recharge the artesian aguifers. In the downward and outward movement from the recharge areas, the water dissolves calcareous material from the aquifers and becomes more alkaline. As the water becomes basic, the iron is precipitated in the aquifers and the concentration of dissolved iron in the water becomes very low. In the discharge areas (fig. 13) the concentrations of dissolved iron are less than 0.1 ppm (fig. 16). Wells tapping aquifers at depths greater than 300 feet below land surface generally produce water with concentrations less than those in fagure 16, at any given location in the county. The concentrations of dissolved iron in water samples collected from all aquifers in Martin County are given in table 4.

Calcium and magnesium (Ca and Mg)

Compounds of calcium and magnesium are common in the limestone, shell beds, and calcareous sands and clays in Martin County. Ground water, especially water of low pH, passing through these sediments will dissolve calcium and magnesium. Most of the calcium and magnesium goes into solution as a bicarbonate salt. The salts of calcium and magnesium are responsible for most of the hardness effect in ground water. Hardness of water is discussed in a separate section. The calcium and magnesium content of water from wells in Martin County is given in table 4.

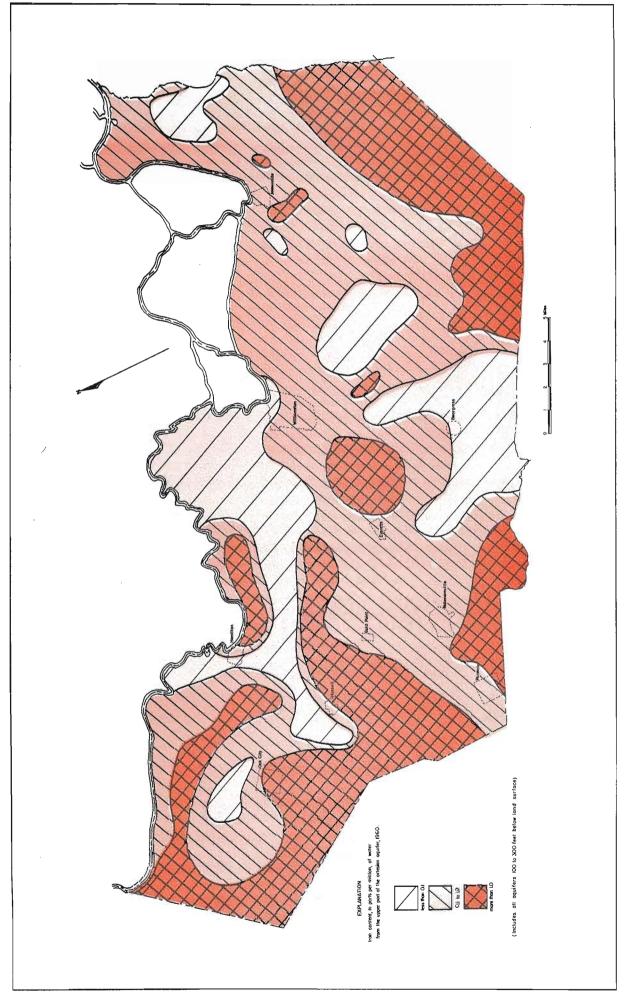
Bicarbonate and carbonate (HCO3 and CO3)

Laboratory analyses of water from wells in Martin County frequently report both the bicarbonate and the carbonate content of water. Field analyses, however, indicate that the bicarbonate type is dominant and the carbonates reported in laboratory analyses generally result from an alteration of the bicarbonates when the water sample is subjected to aeration or to changes in temperature and pressure. The bicarbonate water contains principally cations of calcium and magnesium or of sodium and potassium. Where the cations are calcium and magnesium, the water is usually considered hard but where the cations are principally sodium and potassium the water is very soft. Bicarbonate has little effect on the domestic utilization of water, it may cause boiler foaming or scale deposits in some industrial uses. Zeolite softeners are frequently used on domestic supplies to change calcium and magnesium bicarbonate waters to sodium bicarbonate waters, thus changing hard water to soft water. The bicarbonate content of water from wells in Martin County is given in table 4.

Sulfate (SO_L)

Minerals containing sulfate are frequently present in shell and limestone beds. These minerals, commonly including calcium and nagnesium sulfate, are soluble in ground water and are probably the main source of sulfate ions in the water from artesian wells in Martin County. The sulfate concentrations in water in the county was not high enough to be objectionable for most uses. However, sulfate ions may be reduced by bacterial action to produce hydrogen sulfide and sulfur. Hydrogen sulfide, a gas, has a disagreeable odor and when dissolved in water forms a weak acid which may make the water corrosive.

In the Williamston water system, bacterial action reduces sulfates to sulfides during summer months. The water has a disagreeable odor and corrodes the water system, dissolving some of the iron from iron pipes. This situation could be remedied by the chlorination of the waters to prevent bacterial growth.



IRON CONTENT OF WATER FROM UPPER PART OF ARTESIAN AQUIFERS

Chloride (C1)

More than 90 percent of the dissolved solids in ocean water are chloride salts. The concentration of chloride ions in sea water averages between 19,000 ppm and 20,000 ppm. In ground water in Martin County the chloride content of the water ranges from about 5 ppm to about 1,400 ppm. Most of the chloride is probably derived from the sea water in which the sedimentary formations in the county were deposited. After deposition, the formations were flushed by fresh water to depths of 350 feet to 450 feet below land surface. Generally, all of the aquifers above a depth of 400 feet contain water with less than 250 ppm of chloride. A few water-table wells yield water having concentrations of chloride greater than 250 ppm, indicating that the water has been contaminated by surface pollution. There is no evidence of either lateral or vertical encroachment of salt water into fresh water aquifers in Martin County. There is no economically feasible method of removing chloride from water for domestic purposes. The concentration of chlorides in water from wells is shown in table 4.

Fluoride (F)

Fluoride in ground water is due to the solution of fluoridebearing minerals such as the apatites, the fluorapatites, the phosphates, the micas, and the hornblendes, and organic matter such as shells. Of these, the micas, the phosphates, and the shells are most prevalent in the sediments in this area, and are thought to account for most of the fluoride in the ground water.

Fluoride in concentrations between 1.0 and 1.5 ppm in drinking water aids in reducing tooth decay in children. Fluoride in concentrations in excess of 1.5 ppm may cause permanent mottling of the teeth (dental fluorosis) when used by children (Maier, 1950, p. 1120-1132).

The higher concentrations of fluoride in water in Martin County seem to be restricted to certain geologic formations (see fig. 17). All of the water samples having concentrations of fluoride greater than 1.5 ppm were from aquifers in the Black Creek Formation or along the Black Creek-Peedee Formation contact (see table 4).

Hardness

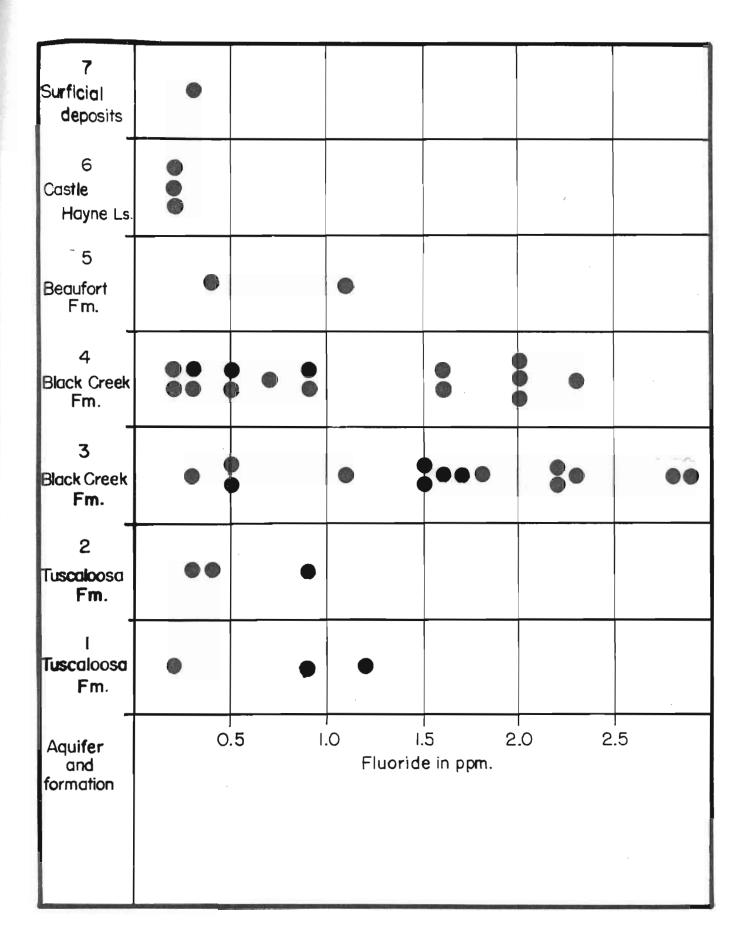
Hardness of water is usually recognized by the increased amount of soap necessary to form and maintain a lather. Hard water is objectionable not only because of its soap-consuming properties but also because it forms scale in boilers and, to a lesser degree, encrustations in cooking utensils. The principal chemical constituents that produce hardness in ground waters are calcium and magnesium. The following classification of water hardness is used by the U. S. Geological Survey:

Hardness as CaCO Classification (ppm)	
0-60 Soft water	
61-120 Moderately hard wa	ater
121-200 Hard water	
200+ Very hard water	

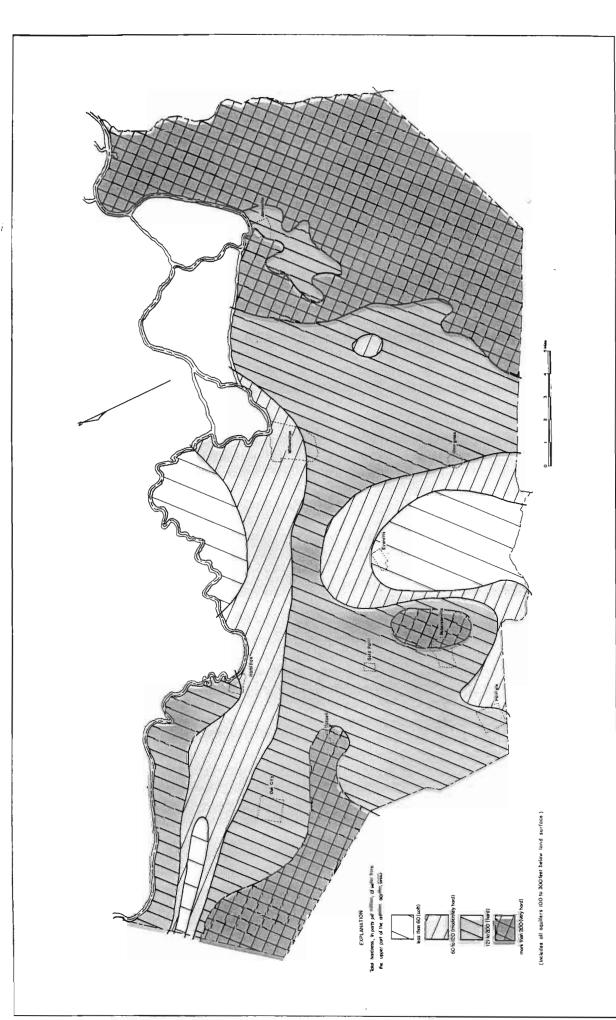
Figure 18 is a map showing the concentrations of hardness in water from wells tapping aquifers from 100 feet to 300 feet below land surface in Martin County. Wells in discharge areas or tapping zones below 300 feet generally produce water of the sodium bicarbonate (soft) type rather than the calcium and magnesium bicarbonate type and do not require treatment for hardness. The hardness of water from samples in Martin County is given in table 4.

Hydrogen-ion concentration (pH)

The hydrogen-ion concentration, expressed as the pH, is a measure of the degree of acidity or alkalinity of the water. The pH of a solution is represented by a number which is the negative logarithm of the concentration of hydrogen ions in moles per liter. Numerically, the pH scale extends from 0-14. A water having a pH value of 7 is assumed to be neutral and the concentration of hydrogen ions is equal to the concentration of the hydroxyl ions. A water with a pH value less than 7 is said to be acid (the concentration of H+ ions is greater than the concentration of OH- ions), and the acidity increases as the numerical pH value decreases toward 0. A water with a pH value greater than 7 is said to be basic (the concentration of OH- ions is greater than the concentration of H+ ions) and the basicity increases as the numerical pH value increases toward 14. Inasmuch as the pH values are the



FLUORIDE CONTENT OF WATER FROM FORMATIONS



HARDNESS OF WATER FROM UPPER PART OF ARTESIAN AQUIFERS

numerical change to the logarithmic base, a water with a pH of 3 is ten times as acid as water with a pH of 4, and conversely a water with a pH of 9 is ten times as basic as a water with a pH of 8.

The pH values given for ground waters are important as an indication of their corrosive potential. Generally speaking, acid waters have a much greater corrosive potential than do the basic waters.

In Martin County water from the water-table aquifer generally has a pH value less than 7.0. The lowest pH measured was 3.4, from a water-table well in an area where the aquifer was contaminated by acid fertilizers. Generally, water from artesian aquifers had a pH value greater than 7.0. The pH of water from wells in Martin County is given in table 4.

CONCLUSIONS

Martin County is underlain by sediments ranging in age from Cretaceous to Recent. The geologic formations from oldest to youngest are the Tuscaloosa, Black Creek, and Peedee Formations of Cretaceous age, the Beaufort Formation of Paleocene age, the Castle Hayne Limestone of middle and late Eocene age, the Yorktown Formation of late Miocene age and the undifferentiated sediments of post Miocene age.

Ground water occurs in seven major aquifers, six of which are artesian and one non-artesian. Massive clay beds of the Yorktown Formation constitute the principal aquiclude in the county. The sediments which underlie the Yorktown Formation contain six major aquifers composed of continuous sand, shell, or limestone beds. The artesian aquifers are separated by clay or silt beds that retard, but do not prohibit, the movement of water from aquifer to aquifer. Aquifers 2 through 6 pinch out along a basal shell bed in the Yorktown Formation which hydrologically interconnects the pinched out ends of the aquifers.

Quantitative studies indicate that most recharge is from precipitation that seeps into the water-table aquifer and, where the water table is higher than the piezometric surface, percolates through the confining layers into the artesian aquifers. Average daily recharge to the artesian aquifers 100 feet to 300 feet deep is calculated to be about 22 million gallons per day, under present hydrologic conditions. The average daily discharge, natural and artificial, is computed to be about 21 million gallons per day. Individual wells will yield 5 to 500 gallons per minute, depending on well construction, well size, and the aquifers tapped by the wells.

Most wells in Martin County could be greatly improved by the use of screen sizes which are designed for use in specific aquifers and by better development of the wells. Municipal and industrial wells should be of the multiple-screen, gravel-pack type to assure high well efficiency and high yield.

Obtaining ground water of good chemical quality may be a frequent problem in Martin County. Generally, wells tapping the water-table aquifer yield water containing concentrations of dissolved iron in excess of 0.5 ppm. Except in recharge areas, artesian water generally has less than 0.3 ppm of dissolved iron. Wells greater than 450 to 500

feet in depth usually produce water with chloride concentrations greater than 250 ppm. Artesian wells less than about 400 feet deep usually produce water with concentrations of chloride ranging from 25 to 50 ppm. The hardness of water from wells tapping shell or limestone aquifers usually exceeds 200 ppm. Generally wells greater than 300 feet deep produce water with hardness of less than 150 ppm.

Wells capable of producing 350 to 450 gallons per minute of water of acceptable chemical quality for most uses may be obtained anywhere in Martin County, if the wells are properly constructed.

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TABLE 3. Record of Wells in Martin County
I. Flanagan
E. Copeland-
A. Everetts
Anthony
L. Mizelle-
A. Everetts
Martin Co. Brd.
Swamplawn Farm
D. Brown-
H. B. Clarkson
ross L. Harrell
Hyman
R. Everetts
T. Winslow-
Martin Co.
Hyman
Wesley, Ayers, and
J. B. Whitfield
Martin Co. Brd.
do
do
Bell-
Keel-
F. E. Manning
)

Yield (gpm) -4.05 -2135 -3.94-3.01-2.73 -6.08 -20.92 -2.55-3.68 -5.80 -11.82 -2.60 -3.62-2.91-28.98 -3.63 Water -1.40 -7.65-1.25-2.31level (ft)-10 Waterbearing casing Depth (ft)97 65 105 132 13 105 105 105 18 165 12 500 24 20 20 0 31 75 17 215 11 Record of Wells in Martin County (Continued) Depth Diam-(in) $1\frac{1}{4}$ eter **6**09 300 18 165 91 105 05 Open end Screen Screen Screen Screen Screen Screen Screen Screen Screen ----qo--- qo---qo---- qo--0**p** ----do-Type of Well P Dug Dug DugDug Dug Dug Dug Dug Dug Dug Educ. Educ. N. C. Forestry Dept. $^{
m o}$ Martin Go. Brd. T. L. Etheridge. Martin Co. Brd. R. H. Salisbury E. R. Bemonson-F. F. Pollard-J. A. Haislip-W. M. Heichmer G. A. Oglesby-P. E. Manning-B. A. Haislip-0wne rW. Alexander-H. K. Blount-J. B. Purvis-Purvis-M. K. Blount-H. R. Purvis-J. S. Ayers-J. Edmonsen-L. L. Lewis-Mr. Hamsey --R. Carson-P. Mathey-J. Dixon-E. North---qp--Ē TABLE 3. 3.7 miles SE of Oak City-5.3 miles NE of Oak City-2,4 miles NNE of Rassellmiles S of Oak City----.. 3 miles E of Oak City-J.5 miles NNE of Hessell-2.4 miles SW of Hassell--5.5 miles SW of Hossell-2 miles W of Hassell-8 miles N of Harsell --1.2 miles N of Parmele-1.4 miles N of Parmele-0.5 miles W of Hassell-3 miles NW of Bassella 0.3 miles N of Parmeleof Hassell---4 miles SW of Hassell-J. I mile E of Hassellof Oak City .0 mile N of Parmelo mile W of Parmele-Location ďΩ miles W lassell-Parmelemiles 1-D0(1-I Down _0Q__ No. Well 47 748 12254595588 40 42 43 中の時 41

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15 22 12 19 150	110	293	114 165 265	470	280	100 218	180	135	140	120	32	06	155	14
do do Screen	qp	qo		qo	Screen	qo		qp	q0	Screen	Bug	Screen	Open end	Dug
61 1.0 miles E of Hassell— H. Banks 62 1.3 miles W of Hamilton— B. Abbitt 64 0.2 miles W of Hamilton— G. Everett— 65 — Bo 65 — Bo 66 2.0 miles W of Roberson—	67 0.2 miles N of Permele G. D. James 68 Parmele		/1	2.0 miles NE of Parmelel.2 miles NW of Beber	0.3 miles W of Hamil ton-	77 Hamilton	79 — Do	81 ——Do		85 Duran Martin Co. Board of Educ.	es E of Hassell P. C. Epstones SSW of Gold	les NW of Roberson-	89 0.9 miles W of Boberson D. Boberson Oville	90 0.4 miles W of Roberson- G. Rowlin-
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TABLE 3. Record of Wells in Martin County (Continued)

Yield (gpm)	1.5 350 350 350 350 0.4 0.2 75 60
Water level (ft)	+4.09 -43.48 -40 -41 -3.23 -4.70 -3.01 +4.45 -6.59 -6.59
Water- bearing zone	せょしょうちょ せてようし てななら みこと ろうみなみ
Depth of casing (ft)	138 155 40 310 320 310 310 100 60 110 115 115 115 115 116 390 390 390 390 390 390
Diam- eter (in)	24 24 24 24 24 24 24 24 24 24 24 24 24 2
Depth	138 155 40 310 320 310 320 310 110 60 60 110 115 115 115 115 115 115 116 390 390 390 390 390 390 390 390 390 390
Type of Well	Screen do-do do Drilled do do do do Screen do do do do Screen do - do - do - do - do - do - - - do - - - - - - - - - - - -
Owne r	C. D. Jenkins——do——do——do——do——do———do———do—————————
Location	0.2 miles W of Roberson- ville- 0.1 miles S of Gold Point A. Bemilton Be Do Bo Bo Bo Book Be 1 mile SE of Gold Point B. 0.5 miles SE of Gold Point B. 0.6 miles S of Gold Point S. 0.4 miles N of Roberson- ville Bo Bo Bo Bo Bo B. 1.5 miles NE of Gold Point B. Robersonville Bo Bo Bo Bo B. 1.5 miles NE of Gold Point B. Robersonville Bo Bo Bo B. 1.5 miles NE of Gold Point W.
Well No.	91 92 93 94 95 95 96 97 96 97 100 100 100 100 110 111 111 111 111 11

0.5 15 3 Yield (mdg) **-0.26** -4.75 -7.81 -2.32 -2.60 -4.60 level -9.02-42.97 -57.12Water -23.35-3.23(ft)bearing Waterzone α \mathbf{C} 4 7 4 5 7 casing Depth (ft)239 15 200 43 320 15 201 325 315 315 315 110 27 $_{
m 0}$ 185 12 180 77 77 80 400 162 90 Record of Wells in Martin County (Continued) Depth Diam- $1\frac{1}{4}$ 24
3
4
24
24 4 1¹/₄ 4 24 4 e ter(in)144,3 5 300 325 315 315 315 35 110 27 112 00^{4} 185 12 180 77 15 239 15 200 43 43 520 15 201 16280 90 Screen -- qo---- qo--Screen --do--Screen Screen Screen Screen Screen -0p--ScreenType of We.11 Dug Dug Dug Educ. -Educ. Educ Co. Martin Co. Board of Martin Co. Bottling Martin Co. Board of Martin Co. Board of W. G. Bullock---4,6 miles SSW of Pweretts H. T. Highsmith A. R. Roberson-M. Hardison-S. Roberson. Roberson Everett. I. L. Murphy-J. R. Everett-Everett 0wner E. Rogerson-Everett T. Parker-Ba. Abbitt-H. Clarks-4.8 miles SSW of Everetts T. Warren-Bunton-Barnhill-B. Wynn-Ι. A 5 TABLE 4.1 miles NNE of Everetts 2.7 miles NNE of Everetts 4.2 miles S of Everetts-3 miles NE of Everetts---3.1 miles Nof Everetts-1.2 miles W of Everetts-0.2 miles N.of Bveretts-5.4 miles NW of Williams-Roberson-Roberson-2 miles ESE of Roberson-2 miles W of Everetts-3 miles N of Everetts-2.4 miles SE of Rober-2.5 miles SE of Reber- $_{
m 0}$ 3.5 miles Siof Location S 1.7 miles sonville sonville⊶ Fveretts---DO---D0villeville---DO---0(I---ville--- Doton165 Well 166 167 No. 162 163 164 143 146 153 154 155 156 157 158 159 160 144 145 147 148 149 150 151 152 161

168 3.6 miles S of Beretts- G. Jones-					+ 01		
1 miles S of Everetts	1 6	<u> </u>	107		12821	89 61	
3.6 miles S of Berrettes	0.9	-2	-20.4 -8.4 -9.3	-3.4	+ .0	-8 -13.1 -7.1 -3.5	
3.6 miles S of Everetts	450	0 22			キキ ~ ~ 5	ンとしなななり	ころろ 4 4 4
10 2.5 miles S of Everetts	206 130 31	25 14 12	108	25 25	21 34	123	165
5.6 miles S of Everetts 6. Jones do	6. 24.	24 24	4 54 54	24	2 24 24 6	24 11 24 24 24 24 24 24 24 24 24 24 24 24 24	74 0 9
2.6 miles S of Everetts— W. A. Mobley W. A. Mobley 1 mile B of Everetts— Nrs. Rogers— Arthur Slade— Williams. On lies W of Williams— On miles W of Beargrass Lester Bryant— On— Do— Do— Do— Do— Do— Do— Do— D	206 130 31	25 14 12	112	25 25	200 200 21 34 137	15 233 236 100 14 400	275 454 290 290 300
3.6 miles S of Everetts- 2.5 miles S of Everetts- Mobilisms S of Everetts- Milliams- Milliams and S of Williams- Milliams and S of Williams- Milliams and S of Williams- More Milliams- More More Milliams- More More Milliams- More More Milliams- More More More More More More More More	do Dug	Dugdo	Drilled Du g Bored	op	Jetted Drilled Dug Bored	Dug Drilleddo Dug Drilled	
	3.6 miles 8 of Everetts- 2.5 miles 8 of Everetts- 1 miles 9 of Everetts-	3.5 miles W of Williams- ton	Williamston————————————————————————————————————	2.2 miles NW of Williams— 2.7 miles W of Williams— ton———————— Joseph Williams———— 3.0 miles W of Williams————————————————————————————————————	2.9 miles SE of Everetts- R. A. Mobley	4.3 miles S of Everetts— John 5.6 miles S of Everetts— W. L. Do W. Do W. Es W.	1.7 miles NW of Beargrass 1.9 miles NW of Beargrass 3.7 miles SW of Williams ton

TABLE 3. Record of Wells in Martin County (Continued)

Ì									
	Location	Owner	Type of Well	Depth	Diam- eter (in)	Depth of casing	Water- bearing zone	Water level (ft)	Yield (gpm)
0,	9.3 miles SW of Williams-					(ft)			
4		Mertin County Home	Drilled	744	4	 	4	 	1 1 1 1
	0(1	Mrs. Van Taylor	qo	280	7	272	, 4	1	
Ä	1.7 miles NW of Williams-	•				•			
ĕ		John Whitleys	qo	278	77	268	7	} 	
3	Williamston	H. E. Harrington	qo	305	23	280	47	-66.24	16
٦.	1.7 miles W center of								
≥	Williamston	T. R. Coltrane	qo	441	4	431	5	-59.25	9
=	Williamston	Town of Williamston	q0	094	12		2	-62	55
1		— is the second of the second	op	450	9	450	2		85
i			qp	450	8,10	450	3,4	-37	200
. 1			qp	644 0	9	044	5	ļ	85
Î			qp	Q*t	9.	740	. 10		85
ř	وم وموزهم 📗 🔾 معناهمارتان فيه ميس بادر نسب فيوا دسترسم تسويدها تعاريب تراها وهو دها بالمواطقة		do	470	20	B 	2		
602	The section and the section of the s		qo	500	18	200	2		300
Ü		Mathison Chemical Co	qo	370	œ		4	+1.3	100
Ē	in the set of the set	Town of Williamston	qo	094	00	094	2		
į.			p	360	50	360	4		
Į.			q0	360	00	360	4		
Ä	1.5 miles & center of								
*	Williamston	Southern Motel	qo	452	4		2		
Ļ	1,6 miles S center of								25
Κi	Williamston	Ress Motel	q0	452	4		2		
Ä	1.7 miles S center of						•		
3	Williamston	Elmo Lilly	Driven	25]]	***************************************	2	7	
∺	1.8 miles S center of				ı				
3	Williamston	Ned Everette	qp	25	$1\frac{1}{4}$		7		
Ø	2.0 miles S center of				·				
≥	Williamston	Rognoke Country Club	Drilled	421	Į	-	2		400
		_!	-	-	-	_	-		

		85			50 0					! !			 	‡ ! !				.		
-41.60 -10.1		-35.55	-34	-35	-24.78 -26.40	-25.07	-30	-33.0			-17.74	-5.03	04-0	C0.02-		741.67		-38.67	-12.88	-21.18
27	2	2	4	וט נו	C 4	7	w 1~	~ r	74	10 L	- 4	7	7 5	C	ıC	5	2	1	īC	5
		423	1	196	14)					154	i	-	310	! ! !		66			-	
7	1	7	7	7	4 01	4	4 6	7	7	27	7	24	4		61	01		7	Ø	5
421	435	423	198	196	300	235	360	28	304	170	315	30	310	190	112	95	441	454	66	150
do Dug	Drilled	qo	qo	qo	Drilled	Drilled	do Driven	Dug Drilled	op	do	Drilled	Bored	Drilled	07	op	Jetted	Drilled	0p	0 p	op
- -	G. F. Harris	John G. Corey	Hubert Biggs	John Hadley	Jack - Bobinson	Beargnass Public School	Rogers Supply Company	yers	J. R. Rogers	Leroy Harrison		76	Bill Harrison	o dreen	Edward A. Hunter	K.B. Kaylor	A. D. Harris	Moose Lodge	Melyid Godard	A. W. Hardinson
center of	liamstonmiles SW of Williams-	miles SW of Williams	5.7 miles SW of Williams-	-	miles N of Beargrass	i				0.8 miles S of Beargrage	S of Beargrass-	miles S of Beargrass-	1 -	S of Williams-	ton	millos S of Millians	mailes S. of Mail in the	miles D or williams	TITES OF OF WILLERING	ton
217 218 219	220	221	222	266		225	227	228 229	230.	231 232		234	- 235	237	258					77

TABLE 3. Record of Wells in Martin County (Continued)

Yield (gpm)								
$\begin{bmatrix} \texttt{Water} & \texttt{Yi} \\ \texttt{level} & (\texttt{ft}) \end{bmatrix}$		-13.15		-7.92				-1.55 -1.25 -1.71 -17.56
Water- bearing zone	2	הייי	V 12	ו טו	U 1U	ו א	ις ις	ろろてててら
Depth of casing (ft)								92
Diam- eter (in)		01 01 47	14	7 -	T	c) o	2) ()	77 77 77 77 77 77 77
Depth	100	90 89 100	22	69	135	100	85	85 103 15 17 31 16 112
Type of Well	Drilled	q0	Driven Drilled	p	new (10 merce)	q0	Driven ~~ado~~~	Drilled Dug Godon Bored Dug Bored Dug
Owner	Mrs. Blanche Hopkins	Noah D, Griffin	Charles E. Edwards	d	ira vonesammamamamamamamamamamamamamamamamamama	H	Arthur D. Coltrain	J. S. Holliday
Location	5,3 miles S of Williams-ton-reserved		miles S o	niles SE of Williams-	10 t	ton	miles S of	ton
Well No.	243	245	248	249	251	253	254	256 257 258 259 260 261

$ \begin{array}{c ccccccccccccccccccccccccccccccccccc$			70 114 80 80 80 80
5 -6.62 5 +2.70 5 -19.96	5	9	24.70 -24.70 -17.14 -17.14 -7.80 -7.80 -7.80 -7.80 -7.80 -7.80
110			102 99 <u>1</u> 99 <u>1</u> 100
7 7 7	114	¢) ()	
105 110 110 110	100	90	45 100 1100 1100 1100 1100 1100 1100 110
op	do Driven	Drilled	Jetted Drilled do do do Jetted Drilled do
262 3.2 miles SE of Beargrass L. J. Peele	267 3.5 miles SE of Williams— ton——————————————————————————————————	1 .	270 5.9 miles B of Beargrass - Hardison - Gurkin - 272 4.8 miles E of Beargrass - Hardison - Griffin - 273 4.5 miles E of Beargrass - Harman L Griffin - 275 4.5 miles E of Beargrass - Harman L Griffin - 275 4.5 miles E of Beargrass - Harman L Griffin - 277 4.7 miles SE of Beargrass Charlie Ann Holliday - 278 5.2 miles SE of Beargrass Charlie Ann Holliday - 279 5.2 miles SE of Beargrass - John A. Revels - 280 5.8 miles E of Beargrass - John A. Revels - 285 5.3 miles E of Beargrass - John A. Revels - 285 5.3 miles E of Beargrass - John A. Revels - 286 5.0 miles E of Beargrass - John A. Revels - 286 5.0 miles E of Beargrass - John A. Hevels - 286 5.

TABLE 3. Record of Wells in Martin County (Continued)

No. Location Owner								1 1						1111	ı
Location			H		1						1				
Location Location Corner Type Depth Diam Depth	Water	level (ft)			1		-6.43					-6.11	-5.17		
Docation	Water-	bearing zone	5	5	5	9	~5	0 IV	5	5	5	7	5	יט יט יט יט	9 .
Location Location Cowner Type Depton	Depth				1	85	128	98				.	1		1
Location Location Cowner Type Depton	Diam-	eter (in)			21	7	42	4	23	7	2	24	2	0	
1.00 1.00	1 0	4	115	110	113	66	20 130	110	130	130	120	20	86	96 100 90 117	115
Location Owner 5.6 miles SE of Williams— 4.8 miles E of Williams— 4.0 miles E of Williams— 4.2 miles E of Williams— 6.0 miles SE of Williams— 7.8 miles SE of Williams— 6.0 miles SE of Williams— 6.3 miles SE of Williams— 6.3 miles SE of Williams— 7.8 miles SE of Williams— 6.3 miles SE of James— 6.3 miles SE of James— 7.8 miles SE of James— 7.9 miles SE of James— 7.1 miles SE of James— 7.2 miles SE of James— 7.2 miles SE of James— 7.3 miles SE of James— 7.4 miles SE of James— 7.5 miles SE of James— 7.7 miles SE of James— 7.7 miles SE of James— 7.7 miles SE of James— 8.7 miles SE of James— 8.8 miles SE of James— 8.9 miles SE of James— 8.0 miles SE of James— 9.1 miles SE of J	Type	of Well	Drilled	0 p	pp	q <u>o</u>	Dug Drilled		qp	q p	q0	Dug	Drilled	qp	q0
Location 5.6 miles SE of William 4.8 miles E of William 4.0 miles E of William 4.2 miles SE of William 6.0 miles SE of Willia 6.7 miles SE of Willia 6.7 miles SE of James 7.8 miles SE of James 7.8 miles SE of James 7.8 miles SE of James 7.6 miles SE of James 9.1 miles SW of James 9.1 miles SW of James	Owner		Havley Hardinson	Clarence Hopkins	L.	Grady Goddard	Martin County	denes Jones	5			Miss Lizzie Taylor	Clarence Hopkins	E. D. Rogers————————————————————————————————————	Thomas Tice
Nell No. No. No. 100 291 292 294 297 298 299 299 300 300 300 300 300 300 300 300 300 3	Location		miles SE of	miles SE of	miles E of	miles E	Out to the second secon	miles Waofamesville miles Waofamesville	ton————————————————————————————————————	miles SE of	miles SW-of	James	SW. of	E of Beargrass-	niles SW of James
<u> </u>	Well	No.	291	292	294	295	296	297	300	301	302	303	304	305 306 307	308

		9	1 1 1		1] 	 				 			1	
	!	+3.83		I I I	-1.12		-4.57	-14.86		-9.90			1			TO COM THE STATE AND STATE
9	9	9	9	9	9	9	9	9	9	9	9	9	9	9	9	9
	en des aus just	95		1	i i	06	89	80	80	80	80		. 06	90		
1 2 2	Day and part and	$1\frac{1}{4}$		61	2	2	C1	c)	61	23	2	13/4	$1\frac{1}{4}$	13/4	23	13/4
100	110	115	125	100	100	100	85	116	115	100	100	110	110	110	110	120
op	0p	qo	op	0 p	qo	0 p	qp	q p	op	0p	qo	qo	qo	op	qo	q 0
	of James- Piney Grove Church		of James-	of James-	of James-	of James-				of Tames—			of Tomos	Milton P. Bennett	of Tomos	
SW SW	8.1 miles SW oville ville SW o	MS.	MS.	5	ville SW o	MS	AS.	AS AS	a la		5 3	MS ASS	8	3		, i
309 310	311	313	314	315	316	317	318	319	750	791	399	323	707	395	706	3

Yield (mds)Water level +7.09 +1,55 +3.54 -12.19+0.50 +7.32 +1.19 (ft)-10.58 Waterbearing zone 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 casing Depth (ft)85 90 09 78 90 90 90 90 90 100 0 Record of Wells in Martin County (Continued) Diam-15/4eter (in)13/4니 무 17 142 O S 20 α Depth 110 110 120110 85 150 110 120 150 900 90 85 86 100 117 Drilled Drilled do Jetted -qo---do----do--op -do--op--dpdo qo--qo--q0do-do-Type of Well Jamesville Hubert & Robert Hardinson of Jamesville Monroe Whittington-W of Jamesville Boy Scout Troop 127 George E. Roberson James Willis Gay-Arthur Roberson Walter Gardnerof Jamesville Lelon Roberson Walter Goddard Noah Roberson Owner L. Manning Gardner Major Gardner David Medlin~ J. H. Perry J. T. Lilly M. Gurkin-W of Jamesville Asa Moore-C. B. TABLE 3. of Jamesville James-James-James-James-James-James⊸ James-James-James Jame s⊸ Jamesof James-0£ $_{0}$ $_{
m ot}$ of SW of Jo $_{
m 0}$ 0 0 $_{
m of}$ J. W-Of Location MS Š 9.8 miles SW ĭ S S_{W} SW SW $\frac{M}{2}$ SW 7.0 miles SW 6.5 miles SM 3.6 miles 8.8 miles 2.9 miles 7.5 miles 8.0 miles 3.5 miles 2.8 miles miles miles miles 1.9 miles 2。4 मां 168 3.8 miles 4.0 miles 4.4 miles ville viller ville ville villevillevillevilleville ville 0(I 2.0 3,2 Well No. 332 334 335 336 342 344 345 327 328 329 331 337 338 339 340 341 343

1/2

2

22

 $\frac{71}{4}$

12

30

FABLE 3. Record of Wells in Martin County (Continued)

Yield (gpm)	
Water level (ft)	-8.88 -23.40 -10.17 -7.50 -7.50
Water- bearing zone	
Depth of casing (ft)	16 80 108 108 100 95 95 95 100 105 105
Diam- eter (in)	1-3-4-4-4-4-4-4-4-4-4-4-4-4-4-4-4-4-4-4-
Depth	168 116 1170 1170 1170 1170 127 127 127 128 128 129 120 120 120 120 120 120 120 120 120 120
Type of Well	Driven Drilleddododo Drilleddo
Owner	Cedar Branch Baptist Ch.— L. H. Lilly———— Milton Griffin———— Edwin Lilly———————————————————————————————————
Location	1.9 miles S of Jamesville 2.5 miles S of Jamesville 2.5 miles S of Jamesville 2.6 miles S of Jamesville 2.8 miles S of Jamesville 2.9 miles S of Jamesville 3.0 miles S of Jamesville 3.0 miles S of Jamesville 4.2 miles S of Jamesville 4.2 miles S of Jamesville 1.2 miles S of Jamesville 1.2 miles E of Jamesville 1.2 miles E of Jamesville 2.0 miles B of Jamesville 2.0 miles B of Jamesville 2.0 miles B of Jamesville 2.9 miles NE of Jamesville 3.8 miles NE of Jamesville 3.8 miles NE of Jamesville
Well No.	377 377 377 377 377 377 381 382 383 383 383 383 393 393 393 393

<u>′</u>										
									500 275	500
-11.52		-10.47								-5.97
9	9	2	9	9	92	9 9	99	9	9	9
140	1 1		175	165 155	300	150	125			
24	62	2 24	01 01	ପପ	25	c) c	J	61 K		<u> </u>
175 27 150	150	180	200	165	190	160	175	190		30
Dug Drilled	qo	do	Drilled	Jetted Drilled	qo	qo	qo	qo	0p	0p
7.0 miles NE of James— Pert Lee Roberson— Do—do— 7.4 miles E of Jamesville C. A. Lilly———— 4.6 miles NE of James—	ville 5.0 miles NE ville 5.0 miles NE	ville———————————————————————————————————	ville 4.6 miles E of Jamesville	4.4 miles E of Jamesville 4.1 miles E of Jamesville 5.8 miles NE of James-		6.5 miles NE of James Arra Sarah Williams	les E of Jamesville Union les E of Jamesville Rev. les SE of James-	SE of	NE of	7.8 miles NE of James.
397 398 399 400	401	404 404	405	400 407 408	409 410	411	412 413 414	415	417	418

TABLE 4. Chemical Analyses of Ground Water From Martin County

b**ear-**ing Zone Water 6.0 6.8 6.8 6.8 6.8 6.7 6.70 7.24 7.24 7.24 7.24 7.30 7.30 8.9 7.24 7.30 8.9 8.9 pHHard-ness as (CaCO₂) Fluo-ride (F) 40 26 8.8 8.8 7.2 7.6 5.4 6.6 77 77 77 5.4 68 7.4 7.8 7.2 7.2 7.2 7.4 5.6 8 8 4 8 8 14.8 Chlo-ride (Cl) (Parts per million) 113 12 7.0 7.0 9.6 9.5 9.5 9.6 9.6 8.4 8.4 8.9 12 12 7 7 $\frac{\mathrm{Sul}}{\mathrm{fate}}$ $(\mathrm{S0}_{4})$ Bicar-bonate (HCO₂) 35 36 36 36 375 371 189 200 200 104 176 189 32 222 222 227 229 229 5.3 9.9 9.8 9.8 9.8 116 117 119 119 119 119 119 119 119 119 ne-sium (Mg) cium (Ca) 0.35 0.01 0.07 0.07 0.04 0.36 0.36 0.36 0.36 0.37 0.10 1.1 Collec-tion Date 8-60 Well No.

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8 7 7 7 7 7 8 9 7 8 8 8 8 8 8 8 8 8 8 8	5.50
129 129 129 131 131 131 136 136 136 137 136 136 136 136 136 136 137 138 138 138 138 138 138 138 138 138 138	51 50 64
1.6	
16 61 62 62 18 268 268 173 16 25 60 49 62 11 12 13 14 11 14 11 14 11 14 11 14 11 14 11 14 16 17 18 18 18 18 18 19 18 18 18 18 18 18 18 18 18 18 18 18 18	2 40 2 16 Survey.
52 7.0 7.0 1.3 86.0 6.0 6.0 6.0 6.0 6.0 6.0 7.7 7.3 8.6 7.3 1.9 6.0 6.0 6.0 6.0 7.3 1.9 1.9 1.9 1.9 1.9 1.9 1.9 1.9	6eological Sur
19 222 286 1117 532 54 559 559 559 579 579 579 579 579 579 579	8.5 S. Geold
22 1.1	8.5 8.5 3ranch, U
21 22 24 15 17 27 28 29 20 21 21 22 23 24 25 26 27 27 27 27 27 27 27 27 27 27	9. 22 12 Wate
0.01 0.40 0.09	0.47 1.80 0.10 Quality of
8	7-60 7-60 7-60 -
23 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	$\begin{vmatrix} 29 \\ 60 \\ 61 \end{vmatrix}$

TABLE 4. Chemical Analyses of Ground Water From Martin County (Continued)

	Water bear- ing Zone	らしてともならのともならとなりもしらっ
٠	Hď	4.71 7.10 4.85 6.9 7.10 7.10 7.10 7.10 7.10 7.10 7.11 6.82 6.70 7.11 8.5 8.5 8.7 7.9 8.7 7.9 8.7 7.9 8.5 8.5 7.9 8.7 7.9 8.7 7.9 8.7 7.9 8.7 7.9 8.7 7.9 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.1 8.7 7.7 8.7 8.7 8.7 8.7 8.7 8.7
	Hard-ness as $\left(\operatorname{CaC}_{0}\right)$	80 61 157 148 195 89 103 17 17 118 174 0.2 98 171 171 148 171 128 37 157 228 135
	Fhuo- ride (F)	
111)	Chlo- ride (Cl)	16 34 277 277 27 14 82 63 63 8.8 8.8 8.2 9.6 50 77 44 25 66 66 66 77 77 77 77 77 77 77
	$\begin{array}{c} \mathrm{Sul}-\\ \mathrm{fate}\\ \left(\mathrm{S0}_{4}\right) \end{array}$	10 25 10.0 10.0 1 4.5 0.6 1.8 2.3 60 20 20 20 20 20 21 14
(rarus F	Bicarbonate ($\mathrm{HC0}_{2}$)	15 8.6 28 3.6 11 5.1 5.1 14 4.4 4.4 9.2 4.7 73 148
	Mag- ne- sium (Mg)	28 28 28 3.6 11 5.1 5.1 24 5.4 2.6 4.4 4.7 20 4.7
	Cal- cium (Ca)	7.9 10 15 52 64 27 48 16 16 16 16 26 26 58 4.9
	Iron (Fe) in Total	1.4 2.4 0.81 0.76 0.76 0.11 1.5 0.28 0.34 0.02 0.02 0.02 0.02 0.03 0.35 0.35 0.35 0.35
	Date of Collec- tion	8-60 8-60 9-8 9-8 9-8 9-8 9-8 9-8 9-8 9-8 9-8 9-8
	Well No.	62 64 65 67 77 77 77 78 80 83 83 85 85

Analyzed by the Quality of Water Branch, U. S. Geological Survey.

TABLE 4. Chemical Analyses of Ground Water From Martin County (Continued)

Water bear- ing Zone	ころころことにはよらなるこうららららはとはならなら
Hď	5.80 7.31 5.00 7.31 7.40 7.20 6.00 7.28 7.28 7.29 7.20 7.20 7.20 8.8
Hard- ness as (CaCO ₂)	59 17 33 18 90 18 203 135 53 228 8 89 102 170 89 170 89 170 18 61 15
Fluo- Fide (C1)	1.6
Chlo- ride (Cl)	2.2 50 16 16 11 13 66 66 66 66 174 33 174 35 103 103 103 12 10 88 88 88 88 88 88 88 86 66
$\begin{array}{c} \operatorname{Sul-} \\ \operatorname{fate} \\ (\operatorname{SO}_{t_{\! 4}}) \end{array}$	18 7 7 52 12 12 1.3 1.3 1.3 1.3 2.2 6.1 6.1 7.0 7.0 7.0 7.0 7.0
Bicar- bonate (HC_0)	52 430 81 332 336 291 39 224 217 73 75 60 60 60 60 60 60 60 60 60 60
Mag- ne- sium (Mg)	18 6.0 6.0 17 10 19 19 19 19 10 10 10 12 12 12 10 10 10 10 10 10 10 10 10 10 10 10 10
Cal- cium (Ca)	4.6 5.2 14 52 71 62 63 24 27 8.9 9.2 8.9 9.2 27 16
Iron (Fe) in Total	0.86 0.3 0.34 1.2 1.2 0.30 1.1 0.36 0.75 1.2 0.02 1.3 1.3 1.3 1.3 1.3 1.3 1.3 1.3
Date of Collec- tion	8 + 60 09 - 7 - 8 - 8 - 60 09 - 7 - 60 09 - 8 - 60 09 - 8 - 60 09 - 7 - 60 09 - 8 - 60 09 - 7 - 60 09 - 8 - 60 09 - 8 - 60 09 - 7 - 8 - 60 09 - 7 - 60 09 - 8 - 60 09 - 7 - 60 09 - 8 - 60 09 - 7 - 60 09 - 8 - 60 09 - 7 - 60 00 - 7 - 60
Well No.	125 125 125 126 127 127 127 127 127 127 127 127 127 127

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7.4.10 6.8.10
158 109 109 109 15 16 17 17 168 62 62 63 136 136 136 136 136 136 136
24 120 120 120 26 26 27 27 27 27 27 27 27 27 27 27 27 27 27
25.2 2.66 2.1 2.1 2.1 2.1 2.1 2.1 2.1 2.1
263 363 374 374 374 374 374 375 376 377 377 377 377 377 377 377
17 12 12 13 13 14 16 16 17 16 17 17 16 17 17 18 18 18 18 18 19 19 19 19 19 19 19 19 19 19 19 19 19
28 29.6 21 17 17 19 19 19 19 19 19 19 19 19 19
0.13 0.04 0.05
8 - 20 - 1 - 20 - 20 - 20 - 20 - 20 - 20
·

Analyzed by the Quality of Water Branch, U. S. Geological Survey.

TABLE 4. Chemical Analyses of Ground Water From Martin County (Continued)

Water bear- ing Zone	4
Hď	8.10 8.4 7.9 8.0 8.1 7.70 7.78 7.58 7.80 7.59 7.99 7.99 7.99 7.99 7.84
Hard- ness as (CaCO ₇)	116 178 54 60 120 148 63 22 22 22 21 14 81 31 10 10 11 16 55 114 57
Fluo- ride (F)	2.2 1.1 1.5 2.0 2.0 2.0 2.3
Chlo- ride (Cl)	8.6 221 275 75 76 9.6 111 3 115 8.6 162 162 177 62 177 62 177 62 177 62 177 62 177 187 197 197 197 198 198 198 198 198 198 198 198
Sul- fate (SO ₄)	24 24 16 16 0.9 2 2 2 145 145 4.4 4.2 13 4.8 3.3
Bicar- bonate (HCO ₃)	365 76 76 76 76 383 383 383 388 368 368 368 368 368 36
Mag- ne- sium (Mg)	20 20 16 16 12 3.0 3.0 3.0 1.3 1.3 1.3 1.3 1.3 1.3 1.3 1.3
Cal- cium (Ca)	13 38 38 21 4.4 4.0 18 6.8 6.8 6.8 1.0 1.1 2.2 1.1 2.2 1.1 1.1 2.8
Iron (Fe) in Total	0.01 0.36 0.1 0.15 0.10 0.98 0.097 0.95 0.95 0.95 0.07 0.07 0.07
Date of Collection	8 8 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9
Well No.	187 188 189 190 191 197 197 198 200 200 200 200 200 200 200 200 200 20

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6.20 7.90 7.50	7.55 7.70 7.55 8.00 7.91 7.91 7.92
63 51 17 7 7 17 8 168 156 17 92 76 17 92 17 34 205 104 17	17 160 208 140 144 148 148 156 136
2.3	
19 71 187 96 63 99 8.8 7 7 7 8.8 9.2 63 14 388 9.6 48 63 63 12 13 13 183	- 183 10 11 11 8 8 8 12 12 8 8 Survey.
22 39 37 30 .7 .5 0.8 0 6.0 6.0 6.0	12 2 7 7 7 7 7 10 10 16 8
22 74 392 316 344 344 42 42 42 43 280 43 247 208	281 374 200 232 216 256 9 203 250
10 .9 .9 11.0 8.5 11 5.9 4.9 17 17 8.0	16 18 12 12 39 8.5 7.0 14 22 Branch, I
14 38 38 1.1 1.1 19 60 60 60 60 60 60 60 60 70 70 70 70 70 70 70 70 70 7	43 60 50 52 10 60 60 30 Water
3.3 10 0.1 0.14 0.05 0.10 0.15 0.15 0.15 0.05 0.10 0.12 0.05 0.13 0.05 0.05 0.13 0.14 0.15 0.17 0.16 0.17 0.17 0.18 0.19	0.71 0.16 0.27 0.05 0.13 0.30 0.23 0.23 0.23
7-60 7-60 7-60 7-60 7-60 8-60 8-60 8-60 7-60 8-60 8-60 8-60 7-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 7-60 8-60 8-60 7-60 8-60 7-60 8-60	٠ ۱
215 215 217 217 227 227 227 227 227 227 227 227	240 241 242 243 244 245 246 246 248 248

Analyzed by the Quality of Water Branch, U. S. Geological Survey.

TABLE 4. Chemical Analyses of Ground Water From Martin County (Continued)

Water bear- ing Zone	うううううううく アイア ううううううう アクラううううう
hф	8.01 8.01 7.3 7.3 7.3 7.3 7.4 8.8 7.4 8.8 7.4 8.8 7.4 8.8 7.4 7.8 7.8 7.8 7.8 7.8 7.8 7.8 7.8
Hard- ness as (CaCO ₂)	140 166 166 176 176 179 179 179 179 179 179 184 184 187 194 187 194 180 180 180
Fluo- ride (F)	
Chlo- ride (Cl)	11277 113777 1138681148681133777 113871188681119
$\begin{array}{c} \operatorname{Sul-} \\ \operatorname{fate} \\ (\operatorname{SO}_{\boldsymbol{4}}) \end{array}$	7.5 9.0 7.0 10 7.0 7.0 7.0 12 7.0 9.0 12 9.0 9.0 9.0 9.0 9.0 9.0
Bicarbonate (HC_0)	264 273 260 273 264 298 298 297 297 298 297 296 296 297 297 297 297 297 297 297 297 297 297
Mag- ne- sium (Mg)	13 17 17 17 18 18 19 19 19 19 11 19 19 19 19 10 10 10 10 10 10 10 10 10 10 10 10 10
Cal- cium (Ca)	50 42 42 52 43 53 45 45 60 60 60 60 77 77 77 77
Iron (Fe) in Total	0.04 0.08 0.08 1.88 0.28 0.30 1.1 0.10 0.25 0.35 0.26 0.07 0.07 0.07 0.07 0.07 0.07 0.07 0.0
Date of Collec- tion	1-60 7-50 7-50 7-50 8-60 7-59 7-59 7-59 7-59 7-59
Well No.	250 251 251 252 253 253 254 254 255 255 255 255 255 255 255 255

ろうらうしょうこううう くらく くらら ららう ららららららららららららら	9
7.4 7.3 7.8 7.4 7.5 6.95 6.95 7.5 7.5 7.7 7.7 7.3 8.2 7.3 8.2 8.2	8.1
160 182 182 184 172 182 173 182 196 196 196 197 197 198 198 199 199 199 199 199 199 199 199	
15 15 16 17 10 10 10 10 10 10 10 10 10 10 10 10 10	
37 00 00 11 12 12 13 14 10 10 10 10 10 10 10 10 10 10 10 10 10	
2008 2008 2008 2008 2008 2009	
26. 27. 29. 29. 29. 29. 29. 29. 29. 29. 29. 29	.
59 473 573 573 574 575 675 675 675 675 675 675 676 677 677	
1.4 0.28 0.18 0.05 0.05 0.05 0.05 0.05 0.05 0.05 0.0	70.0
7	8-60
282 281 282 283 283 283 283 283 283 283 283 283	515

Analyzed by the Quality of Water Branch, U. S. Geological Survey.

Chemical Analyses of Ground Water From Martin County (Continued) TABLE 4.

Water bearing Zone 7.65 7.51 7.50 7.59 7.57 7.38 8. pH $_{(\mathrm{CaCO}_{\mathbf{z}})}^{\mathrm{as}}$ Hardness 214 224 200 228 228 280 Fluo-ride (F) Chlo-ride (Cl) (Parts per million) 8.5 31 9 12 18 9 0.6 3 8.7 9 8 8 Sul-fate (SO₄) 12 14 bonate $(\text{HC}0_{3})$ Bicar-292 300 292 308 348 391 332 389 356 380 380 391 404 404 Magsium (Mg) ne-Cal-cium (Ca) 551 550 550 550 550 660 660 695 695 0.06- $0.12 \\ 0.09$ 0.06 0.17 0.24 0.48 0.98 0.14 0.30 0.30 0.03 0.26 0.42 0.60 0.16 0.13 0.21 0.09 0.22 0.13 0.64 0.76 [otal Iron (Fe) Collec-Date tion09-2 2-60 09-09-7 $_{
m 0}$ Wel1 No.

9979	9 ~	9 [9	9	9	0 9	9	9	9	9		9	9	9	9	2	9	9	9	9	~	9	9	9	9	9	9	9	9	
6.60	7.9	2.00	7.14	-	7.19	(60.)	7.31	7.37	7.40	7.40	9	8.4	9.0	9.5	7.50	6.29				7.21	7.45	7.18	7.28	7.36	7.25	7.11	7.20	7.20	7.00	
292 268 90 216	220 48	336	224 304	284	412	225 236	212	188	240	225	7.40	194	192	180	123	84	214	214	192	176	176	216	224	216	232	256	212	228	232	
				1			1	1		1			.5	1		1		-		-		1	1		1			1		
9.4 10 15	15	18	11	12	10	11.0	10	8.6	27	18	12	13	17	16	<u>.</u> 8.	8.0	9.6	13	9.6	9.5	11	9.5	11	10	7.6	14	8.8	9.6	11	
3.5	4.6 18	7.5	0.7	1.0	1.4	1.1	0,1	2.4	0.6	6.5	5.5	11	1	01	4.5	1	0.6	12	8.0	0.1	0.4	2.0	0.1	0.1	0.1	0.3	0.1	0.1	0.1	
426 360 352	304 38	371 979	379	344	410	408 369	312	256	387	282	584	250	251	237	100	75	274	276	280	568	232	319	240	308	328	378	308	332	348	
30 29 18	35 2.9	77	39	84	09	10	6.6	22	27	23	32	7.9	10	12	5.5	11	38	18	53	0	8	29	0	6.6	0	42	2.7	5.8	56	,
87	52 14	63	58.	35	99	69	69	38	22	47	20	85	4 9	22	53	2.0	45	28	6	20	56.	38	90	20	95	34	80	85	52	
1.19	0.35 2.00	2.57	0.31	1.52	0.20		1.30	•	•	•	•	0.72	1.1	1.5	00.0	50	0.48	0.57	2.20	1.40	1.57	0:38	0.56	0.26	0.30	0.14	0.35	0.18	0.15	
09-8 09-8 09-8	09-8 8-60	09-8	8-60	8-60	8-60	00-2	09-2	09-2	09-2	09-2	09-2	09~2	09-2	09-2	09-2	8~60	09-2	09-2	09-2	2~20	8-60	8-60	09-2	09-2	09-2	8-60	092	09-2	09~2	
346 348 348	350 351	352	354	355	356	358	359	360	361	362	363	364	365	366	367	368	369	220	371	372	373	374	375	326	377	378	379	380	381	,

Analyzed by the Quality of Water Branch, U. S. Geological Survey.

Chemical Analyses of Ground Water From Martin County (Continued) TABLE 4.

	Water bear- ing Zone	000000000000000000000000000000000000000
•	Hd	7.37 7.60 7.60 7.49 7.25 7.25 7.25 7.09 5.60 7.09 7.09 7.09 7.21 6.40 7.29 7.29
	$\begin{array}{c} \text{Hard-} \\ \text{ness} \\ \text{as} \\ (\text{CaCO}_{\overline{2}}) \end{array}$	220 224 264 225 227 225 227 227 220 68 224 224 288 236 288 236 244 132 244 132 288 276 288 288 276 288 276 288 276 288 276 288 276 288 276 277 288 276 276 276 276 276 276 276 276 276 276
	Fluo- ride (F)	
. [Chlo- ride (Cl)	9.2 11 19.4 9.6 9.6 9.6 11 12 12 14 12 14 15 15 15 15
rarus her militani	$\begin{array}{c} \operatorname{Sul-} \\ \operatorname{fate} \\ (\operatorname{S0}_{4}) \end{array}$	0.1 10 10 10 10 10 10 10 10 10 1
(rarus l	Bicarbonate (HCO_3)	336 368 368 368 372 372 374 404 408 388 388 388 388 388 408
	Mag- ne- sium (Mg)	9.9 26 26 27 27 28 28 8.8 8.8 8.8 8.8 8.8 112 24 24 24 26 41 60 60
	Cal- cium (Ca)	72 56 64 64 113 69 75 76 111 76 77 78 78 78 78 78 78 78 78 78
	Iron (Fe) in Total	0.21 0.86 0.15 4.50 0.03 1.1 0.74 0.20 0.00 0
	Date of Collec- tion	7-60 7-60 7-60 7-60 7-60 7-60 7-60 7-60
	.Well	382 384 387 387 387 387 387 387 387 387 400 400 400 400 400 400 400 400 400 40

	9
7.60 7.41 7.50 7.50 7.70 7.55 7.55	!
256 232 212 224 292 292 252 312 137 171	171
	-
14 17 46 9.4 13 11 11 160 188	188
1.5 0.0 0.5 0.5 0.6	S. Geological Survey
332 348 348 376 314 344 366	S. Geolo
29 30 30 32 32 32 36	Branch. U
64 48 75 58 77 87 72 41	'
	944
0.00 0.10 0.75 0.07 0.16 2.8 1.1 6.0	0.1
8-60 0.00 8-60 0.10 8-60 0.75 8-60 0.16 8-60 0.16 8-60 1.1 8-60 0.16 7-60 0.1	the @ual